

LASER GUN SHOCK EXPERIMENTS ON IMPACT VAPOR PLUMES AND ITS IMPLICATION FOR ORIGIN AND EVOLUTION OF PLANETARY ATMOSPHERE. T. Matsui¹, K. Kurosawa¹, S. Ohno¹, T. Kadono², and S. Sugita³, ¹Planetary Exploration Research Center, Chiba Institute of Technology (2-17-1, Tsudanuma, Narashino, Chiba 275-0016, JAPAN), ²School of Medicine, ²Univ. of Occupational and Environmental Health, ³Dept. Complexity Sci. and Eng., Univ. of Tokyo.

Introduction: Atmospheric compositions and pressure are among the most important boundary conditions to investigate the evolution of the surface environment of planets [e.g., 1, 2]. Hypervelocity impacts are thought to play a key role in the origin and evolution of planetary atmospheres. Shock compression/heating and subsequent rapid decompression induce a variety of physical and chemical processes [e.g., 3]. The understanding of physical/chemical behavior of impact vapor plumes is important to investigate the number of geological events [e.g., 4-8]. Such events, however, have not been understood well because of the lack of reliable experimental data on impact-induced vaporization due to experimental difficulties.

In this paper, we describe our experimental approaches for understanding impact-induced vaporization. Recently, high-power lasers used in the studies on nuclear fusion allow us to address extreme conditions on a phase space produced by >10 km/s impacts in a laboratory. We conducted a series of laser gun shock experiments using geologic samples. Then, we applied the results to the geological problems, including atmospheric blow-off on the early Earth and a killing mechanism at the K/Pg impact event.

Experiments: Laser gun shock experiments were carried out using GEKKO-XII HIPER facility at Osaka University. The experimental setup and procedure are described in detail in our previous studies [9-13].

***P-T* Hugoniot measurement for Mg₂SiO₄:** We obtained the Hugoniot curve for forsterite on a pressure – temperature (*P-T*) plane up to 1.2 TPa using the direct drive technique. We captured time-resolved optical signal from a shocked sample using a streaked spectrometer and a VISAR. Figure 1a shows the obtained *P-T* Hugoniot curve for forsterite. The peak shock temperature at >400 GPa is much lower than the M-ANEOS prediction, suggesting that shock-heated forsterite has a higher heat capacity [13, 14]. Such high heat capacity leads to a higher entropy gain, resulting in a higher degree of vaporization after decompression than was previously thought.

Atmospheric blow-off on Earth: The eroded atmospheric mass due to an impact from planets can be calculated based on the *P-T* Hugoniot curve and the sector blow-off model [5]. We investigate the change in the atmospheric pressure on the early Earth during the late veneer phase with a stochastic model [15]. We

found that the pre-existing atmosphere on the early Earth is likely to be in a complete loss during the late veneer phase [14].

Chemical composition of impact vapor plumes:

We conducted the direct measurements of the chemical composition in impact vapor plumes using the flyer acceleration technique. A natural anhydrite (CaSO₄) sample was used as a target to investigate sulfur chemistry in vapor plumes. A quadrupole mass spectrometer (QMS) was used to measure the SO₃/SO₂ molar ratio. Figure 1b shows the SO₃/SO₂ molar ratio as a function of peak shock pressure. We found that SO₃ is the dominant species in impact-induced CaSO₄ vapor plumes at a wide range of peak shock pressure [12].

The K/Pg mass extinction due to acid rain: Our results indicate that a huge amount of SO₃ should be released into the atmosphere after the K/Pg impact. The high SO₃/SO₂ ratio leads to an intense global acid rain rather than global cooling proposed as the killing mechanism in the previous studies [e.g., 12, 15].

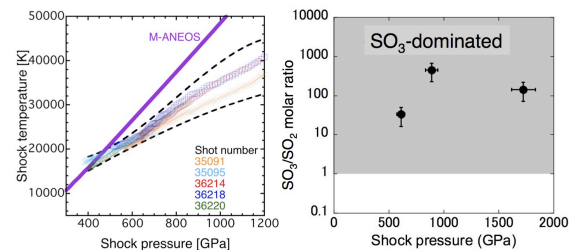


Figure 1. (a) The *P-T* Hugoniot curve for forsterite. (b) The SO₃/SO₂ ratio as a function of shock pressure.

References:

- [1] Matsui and Abe (1986) *Nature*, 319, 303.
- [2] Matsui and Abe (1986) *Nature*, 322, 526.
- [3] Melosh (2007) *MAPS*, 42, 2079. [4] Čuk and Stewart (2012), *Science*, 338, 1047. [5] Vickery and Melosh (1990), *GSA special paper*, 247, 289. [6] Mukhin et al. (1989), *Nature*, 340, 46. [7] Ohno et al. (2004), *EPSL*, 218, 347. [8] Johnson and Melosh (2012), *Icarus*, 217, 416. [9] Kadono et al. (2010), *JGR*, 115, E04003. [10] Kurosawa et al. (2010), *GRL*, 37, L23203. [11] Kurosawa et al. (2012), *JGR*, 117, E04007. [12] Ohno S. et al. (2012), *17th SCCM*, 1426, 851. [13] Kurosawa et al. (2012), *17th SCCM*, 1426, 855. [14] Hicks et al. (2006), *PRL*, 97, 025502. [15] Kurosawa et al. (2013) *44th LPSC*, Abstract #2537.