## LOW INITIAL ABUNDANCE OF <sup>26</sup>AL IN A COMETARY CAI

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**Introduction:** <sup>26</sup>Al is a short-lived radionuclide that decays to <sup>26</sup>Mg with a half-life of 0.717 Myr [1]. It was present in the early solar system and was an important heat source of early accreted planetesimals [2]. The initial distribution of <sup>26</sup>Al in the early solar system remains controversial. Recent analyses of the U-Pb and Al-Mg ages of the andesitic achondrite Erg Chech 002 show that <sup>26</sup>Al was homogeneously distributed within the accretion regions of meteorites [3]. Measured variations in the initial <sup>26</sup>Al/<sup>27</sup>Al of CAIs can be interpreted as either <sup>26</sup>Al variability in the CAI forming region or prolonged (0.4 Myr) thermal processing [4]. The <sup>26</sup>Al distribution in outer solar system objects is not well constrained. Comets, thought to have accreted beyond the orbit of Neptune, contain CAIs at the 2-3% level [5]. Previous measurements of the Al-Mg system in cometary samples (Stardust samples from comet Wild 2, or giant cluster interplanetary dust particles which are probably from comets) show no resolved initial <sup>26</sup>Al/<sup>27</sup>Al [6-9]. All of the previously measured initial <sup>26</sup>Al/<sup>27</sup>Al upper bounds were below the 5.2 × 10<sup>-5</sup> canonical value. Here we report measurements of the <sup>26</sup>Al-<sup>26</sup>Mg system in particle P3-4, a CAI fragment from U2-20 GCP, a giant cluster IDP of likely cometary origin [5].

**Methods:** Initial characterization of the CAI fragment P3-4 TEM was performed at the University of Washington. The particle was embedded in an epoxy cylinder where it was microtomed, creating ~70 nm thick microtome sections and a potted butt. Measurements of the Al-Mg system on P3-4 in the potted butt were made using the Cameca Na-noSIMS 50 and Hyperion plasma source at Washington University in St. Louis. We collected scanning ion images of  $^{24}Mg^+$ ,  $^{25}Mg^+$ ,  $^{26}Mg^+$ ,  $^{27}Al^+$ ,  $^{28}Si^+$  in multi-collection mode using electron multipliers. The mass-resolving power for all species was >5000, sufficient to resolve all interferences. A primary O<sup>-</sup> beam of ~7.5 pA with a spot size of ~200 nm was used to resolve the targeted Al-rich phases from Mg-rich phases surrounding them. During the analyses, the beam was rastered over a 10×10 µm area to include the entire particle in the scanning frame. Prior to the analyses, a larger 20×20 µm area was presputtered to remove the carbon coat.



**Fig. 1.** (Left) Composite RGB element map of the CAI fragment P3-4. (Right) Al-Mg internal isochron plot for P3-4. Dashed red line is canonical, dashed blue line is zero, solid blue line is this measurement.

**Results & Discussions:** The  $({}^{26}\text{Al}/{}^{27}\text{Al})_0$  value obtained from P3-4 is  $(-0.56 \pm 3.12) \times 10^{-5}$  (2 $\sigma$ ). The 2 $\sigma$  upper bound for initial  ${}^{26}\text{Al}/{}^{27}\text{Al}$  is  $2.56 \times 10^{-5}$ . This upper bound corresponds to about 1 Myr after CAI formation. Similar to other cometary particles, P3-4 shows no evidence for the incorporation of live  ${}^{26}\text{Al}$  during its formation or a subsequent heating event. There are three possible explanations for lower than canonical  ${}^{26}\text{Al}$  in P3-4 and other cometary particles: they may have formed (or were altered) 1) before  ${}^{26}\text{Al}$  was injected into the solar nebula, 2) in a region where  ${}^{26}\text{Al}$  was less abundant, or 3) after  ${}^{26}\text{Al}$  decayed.

**References:** [1] Norris T. L. et al. (1983) Journal of Geophysical Research: Solid Earth 88.S01: B331-B333. [2] Dauphas N. and Chaussidon M. (2011) Annual Review of Earth and Planetary Sciences 39:351-386. [3] Reger P. M. et al. (2023) Geochimica et Cosmochimica Acta 343: 33-48. [4] Kawasaki N. et al. (2020) Geochimica et Cosmochimica Acta 279: 1-15. [5] Joswiak D. J. et al. (2017) Meteoritics & Planetary Science 52.8:1612-1648. [6] Ishii H. A. et al. 41st Annual Lunar and Planetary Science Conference. No. 1533. 2010. [7] Nakashima D. et al. (2015) Earth and Planetary Science Letters 410: 54-61. [8] Ogliore R. C. et al. (2012) The Astrophysical Journal Letters 745.2: L19. [9] Ogliore R. C. et al. (2020) Geochimica et Cosmochimica Acta 271: 116-131.