

**UREILITES: MIXES OF A VESTA-LIKE PARENT BODY AND AN IMPACTOR FROM PROTO-MARS.**

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**Ureilites:** Ureilites are unusual and enigmatic meteorites. They are as carbon-rich (~3 wt% C [1]) as carbonaceous chondrites, but compositionally resemble ordinary chondrites [2]. Their <sup>50</sup>Ti and <sup>54</sup>Cr abundances place them in the inner solar system [3], near 2.7 AU [4]. They are achondrites but saw only partial (~15%) extraction of melt [2]. Most (95%) ureilites are monomict with olivines of uniform Mg# varying from 74 to 96, strongly peaked at 80. Several properties (e.g.,  $\Delta^{17}\text{O}$  and  $\delta^{18}\text{O}$ , Fe/Mn ratios), correlate with Mg# [5]. The range of Mg# has been attributed to “smelting”:  $(\text{Mg,Fe})_2\text{SiO}_4 + \text{CaO (melt)} + \text{SiO}_2 \text{ (melt)} + \text{C} \rightarrow (\text{Mg,Ca})\text{SiO}_3 + \text{Fe} + \text{CO (gas)}$ . The equilibrium Mg# depends on pressure and depth in the ureilite parent body (UPB) [6,7], but why ureilites would so preferentially sample Mg# of 80 is a mystery [8]. C content does not correlate with Mg# [9]. Pyroxene thermometers suggest the UPB last equilibrated at about 1050-1100°C [10], after peak temperatures  $\approx 1200\text{-}1300^\circ\text{C}$  [2,11,12]. Thermal models including melt migration reproduce these temperatures if the UPB had radius  $\approx 100\text{-}250$  km and formed at  $t=0.6$  Myr (relative to CAIs) [13]. Soon after  $t=5$  Myr [14], the UPB was catastrophically disrupted by impact at  $\sim 5$  km/s (consistent with typical S4 shock stages; [15]), as inferred from reduction rims around olivines attributed to “smelting” initiated by release of pressure by an unroofing event, followed by a quench in temperatures. The UPB must have broken into chunks  $< 10$  m in size [16] that reassembled into ureilite daughter bodies (UDBs) [2,17], from which ureilites derive. Impact shock would have produced the copious nanodiamonds in ureilites [18], but this does not explain the large (100  $\mu\text{m}$ ) single-crystal diamonds with  $\delta^{15}\text{N}$  zoning observed in the polymict ureilite Almhata Sitta MS-170, which formed in metallic melt in a planetary mantle at pressures  $\approx 4$  GPa [19,20].

**Model for Ureilite Origins:** We hypothesize the following. The UPB formed at 2.7 AU at  $t=0.6$  Myr. Disk models predict it was 1.3wt% CAIs [21]. Its initial composition was like 0.63 H+0.33 CV+0.04 CI chondrites, similar to the 0.75 H+0.25 CV mix inferred for Vesta [22]. This composition yields Mg# 80 olivines and  $\sim 1\text{wt}\%$  C. Some silicates at low pressures underwent equilibrium smelting, forming high-Mg# olivines. The smelted fraction constrains the peak pressure; we infer the UPB radius was 173 km. We interpret the  $\approx 25\%$  of olivines with Mg#  $< 80$  to be from the impactor, mixed in during the impact. Its composition corresponds to end-member “A” of [5]. The impactor had 30% the mass of the UPB, consistent with [8], and had radius  $\approx 115$  km. It delivered metal and abundant C, including large diamonds. During reassembly into UDBs, temperatures were  $\approx 1100^\circ\text{C}$ , so that silicates did not melt, but metallic melts from both bodies mixed, explaining HSE abundance trends [23]. Carbon was redistributed by the melt.

**Largest Daughter Body:** We further hypothesize that 15 Eunomia at 2.64 AU is the largest UDB. Based on the radii above and the 5 km/s impact speed, using [24] we estimate a radius of the largest UDB  $\approx 139$  km; Eunomia's radius is 132 km. Although the polymict ureilite Almahata Sitta derived from the F-type asteroid 2008 TC<sub>3</sub> [25], most ureilites are spectrally associated with S-type asteroids [26]; Eunomia is the largest S-type asteroid. 15 Eunomia has an extensive and ancient collisional family with a dynamical pathway to deliver fragments to 2.55 AU [27]. We suggest the F-type asteroid 438 Zeuxo at 2.55 AU is from the Eunomian family, and 2008 TC<sub>3</sub> derived from it and underwent drifted to the 3:1 resonance at 2.5 AU to reach Earth, consistent with its inferred dynamics [25].

**Impactor Origin:** Finally, we suggest the impactor derived from the proto-Martian surface at  $t\sim 5$  Myr, after magma ocean crystallization but before mantle overturn [28]. It would be Fe-rich bulk Mars, plus late-accreted carbonaceous chondrite material, with Mg#  $\approx 74\text{-}80$ . We find in oxygen isotopes, Mg#, and Fe/Mn it would match end-member “A” of [5]. The compositional similarity between the UPB (Mg# = 80) and the impactor is somewhat coincidental, but they did derive from similar starting materials. Ejection by the Borealis basin impactor would have generated  $> 100$  fragments larger than 100 km in radius that would have impacted objects out to 2.9 AU at 5 km/s [29]. The impactor could have delivered diamonds, formed in Mars's mantle at  $P > 4$  GPa, to the UPB and UDBs.

**References:** [1] Hudon, P et al. (2004) *LPSC* 35, 2075. [2] Goodrich, C et al. (2015) *MAPS* 50, 782. [3] Warren, P (2011) *GCA* 75, 6912. [4] Yamakawa, A et al. (2010) *ApJ* 720, 150. [5] Barrat, J et al. (2017) *EPSL* 478, 143. [6] Walker, D & Grove, T (1993) *Metics.* 28, 629. [7] Singletary, S & Grove, T (2003) *MAPS* 38, 95. [8] Michel, P et al. (2015) *P&SS* 107, 24. [9] Goodrich et al. 1992. [10] reference. [11] Goodrich, C (2004) *ChEG* 64, 283. [12] Goodrich et al. 2007. [13] Wilson, L et al. (2008) *GCA* 72, 6154. [14] Goodrich, C et al. (2010) *EPSL* 295, 531. [15] Stoeffler et al. (1991) *GCA* 55, 3845. [16] Herrin, J et al. (2010) *MAPS* 45, 1789. [17] Downes, H et al. (2008) *GCA* 72, 4825. [18] Lipschutz, M (1964) *Science* 143, 1431. [19] Miyahara, M et al. (2015) *GCA* 163, 14. [20] Nabiei, F et al. (2018) *Nat Comm* 9, 1327. [21] Desch, S et al. (2018) *ApJS* 238, 11. [22] Righter, K & Drake, M (1997) *MAPS* 32, 929. [23] Goodrich, C & Desch, S (2019) this conference. [24] Leinhardt, Z & Stewart, S (2012) *ApJ* 745, 79 [25] Jenniskens, P et al. (2010) *MAPS* 45, 1590. [26] Gaffey, M et al. (1993) *Icarus* 106, 573. [27] Carruba, V et al. (2007) *A&A* 473, 967. [28] Elkins-Tanton, L et al. (2003) *MAPS* 38, 1753. [29] Hyodo, R & Genda, H (2018) *ApJ* 856, L36.