

EXPERIMENTAL INVESTIGATION OF FRACTIONATION DURING SOLIDIFICATION OF AN INCOMPATIBLE-ELEMENT-RICH LUNAR BASALT FROM APOLLO 12.

A. C. Stadermann¹, B. L. Jolliff², M. J. Krawczynski², and C. W. Hamilton¹, ¹University of Arizona, Lunar and Planetary Laboratory (1629 E University Blvd, Tucson, AZ 85721 USA; acs@lpl.arizona.edu), ²Washington University in St. Louis, Department of Earth and Planetary Science, McDonnell Center for Space Sciences (1 Brookings Drive CB 1169, Saint Louis, MO 63130 USA).

Introduction: Apollo 12 sample 12032,366-18 is a basaltic rock fragment with KREEP enrichment (K, REE, P, and other incompatible elements) [1,2], including mesostasis regions that may have formed in part by silicate-liquid immiscibility (SLI). The mesostasis of 12032,366-18 is texturally unlike other samples that exhibit SLI; the high Si and K, and the inclusion of high-Z phases make this mesostasis occurrence unusual.

In this work, we seek to understand the phase relationships and constrain the processes that led to formation of mesostasis in 12032,366-18 through a series of crystallization experiments designed to reproduce the liquid line of

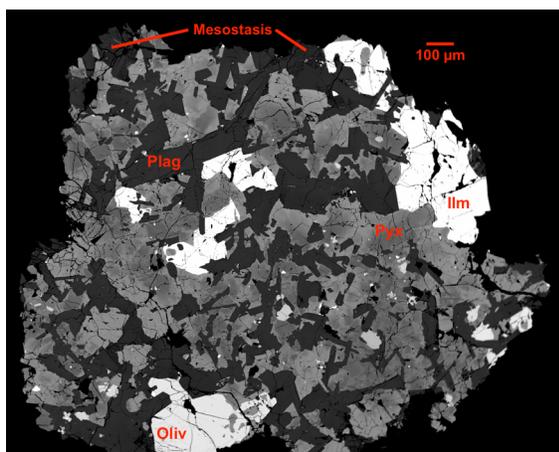


Figure 1: Backscattered-electron (BSE) image of 12032,366-1. Note zoning in pyroxene and the several mesostasis regions at the top of image.

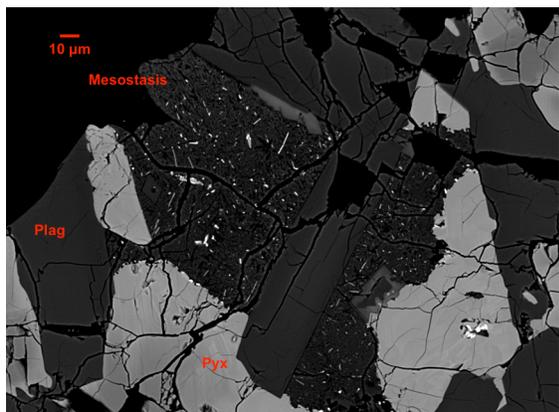


Figure 2: BSE image of mesostasis region in the top left of Fig. 1. The mesostasis includes SiO₂, Ba-K-feldspar, Fe-Ti oxide, zircon, and Ca-phosphate.

descent for the sample's bulk composition. Because of its small size and possibly non-representative modal mineralogy (Fig. 1) experiments were designed to investigate whether the observed bulk composition would reproduce the existing assemblage, including the mesostasis.

Results and Discussion: Sample 12032,366-18 contains high- and low-Ca pyroxene zoned to Fe-rich compositions, plagioclase, Fe-rich olivine, and ilmenite. Several regions of mesostasis have a fine-grained matrix rich in Si and K. The mesostasis has ~80 wt% SiO₂, which is unlikely to have formed by fractional crystallization.

A experimental series at 1-atm pressure and lunar-like fO₂ (iron-wüstite) was conducted to explore fractional crystallization systematics. The mineralogy and chemistry of precipitated phases closely match those observed in the lunar sample. SLI was encountered at temperatures below 1000°C, and at 985°C, SLI was extensive. An example of the compositions of coexisting Fe-rich and Si-rich melt phases is given in Table 1. The Si-melt is not as rich in Si as the mesostasis in 12032,366-18, but the Si content of the Si-melt increases as temperature of the experiment decreases. Merrillite crystallization occurred after the onset of SLI.

Conclusions: Low-P experiments support crystallization of 12032,366-18 from a liquid similar to its analytical-determined bulk composition, including the occurrence of Fe-rich olivine and ilmenite. Fractional crystallization of the major mineral assemblage led to late-stage SLI beginning ~1000°C, which can explain the occurrence of the mesostasis in 12032,366-18.

Acknowledgements: This work was funded by NASA grant NNX15AJ25G, and SSW NNX15AL60G.

References: [1] Jolliff et al. (2005) *Lun. Plan. Sci.* 36, #2357; [2] Barra et al. (2006) *GCA* 70, 6016-6031.

Table 1: Average compositions of Fe-melt and Si-melt at 985°C. Zr is concentrated in the Fe-rich melt, while the K is concentrated in the Si-rich melt.

	SiO ₂	TiO ₂	Al ₂ O ₃	FeO	MnO	MgO	CaO	BaO	Na ₂ O	K ₂ O	P ₂ O ₅	ZrO ₂
Fe-melt	46.22	2.70	7.37	23.54	0.30	1.27	11.39	0.47	0.55	1.32	2.07	0.92
Si-melt	67.05	0.94	11.62	8.27	0.09	0.18	3.79	0.45	0.85	5.36	0.37	0.32