

THE RANGE OF INITIAL $^{10}\text{Be}/^9\text{Be}$ RATIOS IN THE EARLY SOLAR SYSTEM: A RE-ASSESSMENT BASED ON ANALYSES OF NEW CAIs AND MELILITE COMPOSITION GLASS STANDARDS.

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Introduction: Beryllium-10, which decays to ^{10}B with a half-life of 1.4 Ma, is produced almost exclusively by irradiation reactions induced by solar or galactic cosmic rays ([1], and references therein). Previous studies have demonstrated that primitive calcium-aluminum-rich inclusions (CAIs) in several carbonaceous chondrite types record the prior existence of this short-lived radionuclide with an initial $^{10}\text{Be}/^9\text{Be}$ ratio typically in the range of ~ 3 to $\sim 10 \times 10^{-4}$ [2-8]. These secondary ion mass spectrometer (SIMS) studies used NIST glass standards to calculate the Be/B ratios in the samples analyzed. It was shown recently that for an accurate quantification of the Be/B ratios in melilite via SIMS analyses, the relative sensitivity factor (RSF, defined as $(^9\text{Be}^+/^{11}\text{B}^+)/(^9\text{Be}/^{11}\text{B})$) needs to be determined using melilite composition glass standards [9]. In this study, we synthesized and determined the composition of melilite glass standards to revise our previously obtained data for several CAIs [10], and to accurately quantify the initial $^{10}\text{Be}/^9\text{Be}$ ratio of a new CAI (Saguaro) from Northwest Africa (NWA) 5508 (CV3) and the FUN CAI CMS-1 [11, 12] from Allende (CV3). The goal of this work is to obtain better constraints on the origin of ^{10}Be in these refractory inclusions and, ultimately, the astrophysical birth environment of our Solar System.

Analytical Methods: Mineralogical characterization of the Saguaro CAI was performed using the JEOL JXA-8530F electron microprobe at ASU, while that for CMS-1 (including element maps and mineral compositions) was reported previously by [11]. Three melilite-composition glasses doped with varying amounts of Be and B were synthesized using procedures described previously by [12]. The major element compositions and compositional uniformity of these glasses were determined using the JEOL JXA-8530F electron microprobe at Arizona State University (ASU), Be and B concentrations were measured using the iCAP-Q ICP-MS at ASU, and the RSF for accurately determining the Be/B ratios was determined for the Cameca IMS-6f SIMS at ASU and the Cameca IMS-1290 SIMS at UCLA. The Be-B isotope systematics of the Saguaro CAI were measured on the Cameca IMS-6f SIMS using techniques described by [12]; a primary beam current of 15-30 nA resulted in a ~ 30 μm diameter spot. Analyses of the B isotopic composition and the Be/B ratios of individual melilites in the CMS-1 FUN CAI were performed with the Cameca IMS-1290 SIMS at UCLA using an $^{16}\text{O}^-$ primary beam generated by a *Hyperion-II* source. The primary beam current was 10-15 nA (resulting in a beam size $\sim 5 \times 4$ μm), the mass resolving power was ~ 2000 , and ^9Be (10s), ^{10}B (20s), ^{11}B (15s), and $^{28}\text{Si}^{++}$ (3s) were counted for 10-50 cycles depending on the B count rates. IMt-1 clay and NIST614 glass standard were measured during the analysis sessions on both SIMS instruments to correct the measured $^{10}\text{B}/^{11}\text{B}$ ratio for instrumental mass fractionation.

Results and Discussion: The three synthetic melilite composition glasses have Åk contents of 20, 54, 75 and $^9\text{Be}/^{11}\text{B}$ atomic ratios of ~ 11 , 21, and 56. The resulting RSFs are 1.9 ± 0.2 and 2.1 ± 0.2 (2σ) for the IMS-6f and the IMS-1290, respectively; these are $\sim 30\%$ lower than RSFs calculated for non-matrix matched NIST glasses for these instruments. We recalculated the initial $^{10}\text{Be}/^9\text{Be}$ ratios based on our previous data [10] using the new RSF, and these recalculated values are in the range of $(6-50) \times 10^{-4}$ (somewhat broader than that determined by [9] of $(9-20) \times 10^{-4}$). The CAI Saguaro from NWA 5508 (CV3) is a large (~ 18 mm on the longest axis) coarse-grained compact type A inclusion. Our Be-B data for this inclusion (obtained with the IMS-6f) yielded a $^{10}\text{Be}/^9\text{Be} = 7.9 \pm 2.0 \times 10^{-4}$ and an initial $^{10}\text{B}/^{11}\text{B} = 0.239 \pm 0.002$ (2σ). We had previously measured the Be-B isotope systematics in the FUN CAI CMS-1 with the IMS-6f at ASU but had been unable to locate clean melilite areas with Be/B ratios $> \sim 5$ [12]. To obtain better constraints on the initial $^{10}\text{Be}/^9\text{Be}$ ratio for this sample, we re-analyzed it with the IMS-1290 at UCLA (which is equipped with a new *Hyperion* source) because a smaller spot size could be achieved with this instrument. Seven analysis spots on CMS-1 melilites range in $^9\text{Be}/^{11}\text{B}$ ratios from ~ 1 to ~ 87 . However, even with these new Be-B data, we were only able to define an upper limit on the $^{10}\text{Be}/^9\text{Be}$ ratio for this CAI of $\leq 3.8 \times 10^{-4}$, with an initial $^{10}\text{B}/^{11}\text{B} = 0.262 \pm 0.005$ (2σ , which is $\sim 5\%$ higher than the chondritic value 0.248 [13]). This upper limit is slightly higher than the $^{10}\text{Be}/^9\text{Be}$ ratio of $\sim 3 \times 10^{-4}$ defined by three other FUN CAIs [4,7,14] but is still lower than the range for normal CAIs. These results further suggest that FUN inclusions were formed in a different irradiation environment than normal CAIs, and that ^{10}Be is predominately produced by spallation of CAI precursors in the solar nebula.

References: [1] Davis A.M. & McKeegan K.D. (2014), *Treatise on Geochemistry* (2nd Ed.), p.361. [2] McKeegan K. et al. (2000) *Science* 289, 1334. [3] Sugiura N. et al. (2001) *MAPS* 36, 1397. [4] MacPherson G. et al. (2003) *GCA* 67, 3165. [5] Chaussidon M. et al. (2006) *GCA* 70, 224. [6] Liu M.-C. et al. (2009) *GCA* 73, 5051. [7] Wielandt D. et al. (2012) *ApJ* 748, 25. [8] Gounelle M. et al. (2013) *ApJ* 763, 33. [9] Fukuda K. et al. (2016) *MetSoc* 79 abstract #6459. [10] Dunham E. et al. (2016) *MetSoc* 79, abstract #6222. [11] Williams C.D. et al. (2017) *GCA* 201, 25. [12] Dunham E. et al. (2017) *LPS* 48 abstract #1507. [13] Zhai M. et al. (1996) *GCA* 60, 4877. [14] M.-C. Liu et al. (2017) *LPS* 48 abstract #1249.