

**Micro-XANES Determination on the Oxidation States of V and Fe  
in Olivine-Hosted Glass Inclusion and Groundmass Glass of Yamato 980459**

R. Nakada<sup>1</sup>, T. Usui<sup>2</sup>, M. Ushioda<sup>3</sup>, and Y. Takahashi<sup>4</sup>

<sup>1</sup> Kochi Institute for Core Sample Research, JAMSTEC, Monobe 200, Nankoku, Kochi 783-8502, Japan, <sup>2</sup> Earth-Life Science Institute, Tokyo Institute of Technology, Meguro, Tokyo 152-8550, Japan, <sup>3</sup> Geological Survey of Japan, AIST, Central 7, 1-1-1 Higashi, Tsukuba, Ibaraki 305-8567, Japan, <sup>4</sup> Department of Earth and Planetary Science, The University of Tokyo, Hongo 7-3-1, Bunkyo, Tokyo 113-0033, Japan.

**Introduction:** The redox condition, or oxygen fugacity ( $fO_2$ ), of a magma controls the stability and composition of crystallizing and volatile phases in a magma, and provides information on the genesis, differentiation, and a source region of the mantle. A knowledge on the evolution of the Martian interior has been acquired by the analysis of the Martian meteorites. Mineralogical oxybarometers such as Fe-Ti oxides have been conventionally used to estimate the  $fO_2$  of Martian magmas (e.g., [1]–[5]). Partitioning of redox sensitive elements has been more recently used to evaluate the  $fO_2$ , combined with laboratory experiments (e.g., [6]–[10]). These previous studies, however, did not pay much attention to olivine-hosted glass inclusion and groundmass glass that generally crystallized at the earliest and the latest stages of the formation of basaltic rocks, respectively; investigation of these contrasting phases will provide the information on the transition of  $fO_2$  condition during the basalt magmatism. Therefore, this study performed micro X-ray absorption near edge structure ( $\mu$ -XANES) analysis on V and Fe in both inclusion and groundmass glasses of a Martian meteorite (Y98: Yamato 980459) to estimate the  $fO_2$  condition of Martian magma.

**Sample:** Y98 is an olivine-phyric shergottite, displaying a porphyritic texture with olivine megacrysts. Since Y98 represents a primary melt composition that evolved in a close-system ([11], [12]), Y98 is an ideal meteorite sample to compare the valence states of V and Fe in the inclusions and groundmass glass.

**Analysis:** A polished thin section of Y98 (#51-2) was used in this study. The V K-edge (5465 eV)  $\mu$ -XANES analyses were performed using the bending-magnet beamline 10.3.2 at the Advanced Light Source (ALS), Lawrence Berkeley National Laboratory, California. The X-ray beam was focused using a K-B mirror to a final spot size of 3.3 (vertical)  $\times$  2.8 (horizontal)  $\mu\text{m}^2$ . The Fe K-edge (7111 eV)  $\mu$ -XANES spectra were measured at the BL-4A of Photon Factory (PF; Tsukuba, Japan), where the X-ray beam was focused using a K-B mirror to a final spot size of 4.5 (V)  $\times$  4.0 (H)  $\mu\text{m}^2$ . Prior to the  $\mu$ -XANES measurements, X-ray fluorescence (XRF) mapping which was scanned in 5  $\mu\text{m}$  step were obtained to determine the analytical spot with referring to the backscattered electron images [12]. The oxidation state analyses on V and Fe were performed by analyzing the pre-edge features ([13], [14]).

**Result and Discussion:** The mean V valence varied from 3.11 to 3.34 for the olivine-hosted glass inclusions, whereas 3.28 to 3.65 for those of the groundmass glasses. The oxidation state of Fe is strongly correlated with that of V, and varied from 1.85 to 2.01 for the inclusions and 1.90 to 2.36 for the groundmass glasses. The mean  $fO_2$  values relative to the iron-wüstite (IW) buffer for the inclusion glass was IW-0.1, whereas that of groundmass glass was IW+0.8. Previous studies reported similar redox conditions. The oxygen fugacity during the spinel crystallization of Y98 was estimated to be IW+0.9 [15]. The partitioning of V between olivine and melt suggested that the  $fO_2$  of Y98 was IW+0.9 [7]. Chemical composition of olivine megacryst suggested  $\sim$ IW+1 [12]. The partitioning of V and Cr between pyroxene and melt was IW $\pm$ 0 to IW+1 [8]. The current study based on direct V and Fe  $\mu$ -XANES analyses on glass inclusion and groundmass glass showed consistent  $fO_2$  values with previous studies. In addition, it should be noted that the current study successfully indicated the redox evolution during igneous process occurred in Mars.

**References:** [1] Stolper E. and McSween Jr. H.Y. 1979. *Geochimica et Cosmochimica Acta* 43:1475–1498. [2] Steele I. M. and Smith J. V. 1982. *Journal of Geophysical Research* 87:A375–A384. [3] McSween Jr. H. Y. et al. 1996. *Geochimica et Cosmochimica Acta* 22:4563–4569. [4] Ghosal S. et al. 1998. *Contributions to Mineralogy and Petrology* 130:346–357. [5] Herd C. D. K. et al. 2001. *American Mineralogist* 86:1015–1024. [6] Wadhwa M. 2001. *Science* 291:1527–1530. [7] Shearer C. K. et al. 2006. *American Mineralogist* 91:1657–1663. [8] Karner J. M. et al. 2007. *American Mineralogist* 92:1238–1241. [9] McCanta M. C. et al. 2009. *Meteoritics & Planetary Science* 44:725–745. [10] Papike J. J. et al. 2013. *American Mineralogist* 98:2193–2196. [11] Greshake A. et al. 2004. *Geochimica et Cosmochimica Acta* 68:2359–2377. [12] Usui T. et al. 2008. *Geochimica et Cosmochimica Acta* 72:1711–1730. [13] Sutton S. R. et al. 2005. *Geochimica et Cosmochimica Acta* 69:2333–2348. [14] Wilke M. et al. 2001. *American Mineralogist* 86:714–730. [15] McKay G. et al. 2004. Abstract #2154. 35th Lunar & Planetary Science Conference.