

1:2500 GEOMORPHOLOGICAL MAP OF THE INTERCRATER REGION BETWEEN SHACKLETON CRATER AND SHOEMAKER CRATER OF THE LUNAR SOUTH POLE. F. B. Wróblewski¹, K. R. Frizzell², G. R. L. Kodikara³, M. Kopp⁴, K. M. Luchsinger⁵, A. Madera², M. L. Meier¹, T. G. Paladino⁶, R. V. Patterson⁷, C. J. Tai Udovicic⁸, and D. A. Kring^{9,10}. ¹University of Idaho (Email: frankw@uidaho.edu), ²Rutgers University, ³University of Wisconsin-Milwaukee, ⁴Boston College, ⁵New Mexico State University, ⁶Idaho State University, ⁷University of Houston, ⁸Northern Arizona University, ⁹Lunar and Planetary Institute, Universities Space Research Association, ¹⁰NASA Solar System Exploration Research Virtual Institute.

Introduction: Shackleton crater (21 km diameter, 4 km deep [1], $3.43 \pm 0.04/-0.05$ Ga [2], $131^\circ\text{E } 89.6^\circ\text{S}$) is a deep Late Imbrian impact crater at the lunar south pole along the margin of the South Pole-Aitken (SPA) basin [1]. Shackleton is separated from Shoemaker crater (50 km diameter, 2.1 km deep [1], $4.15 \pm 0.02/-0.02$ Ga [3], $50^\circ\text{E } 88^\circ\text{S}$) by an intercrater plain. Both Shackleton and Shoemaker craters host large permanently shadowed regions (PSRs) which contain volatile elements [4,5] that may be used to trace the evolution of volatiles on the Moon and provide *in situ* resources [6,7]. In addition, the rims of both craters are illuminated topographic highs feasible for solar power [8,9]. Despite those favorable conditions, crater wall slopes $>25^\circ$ make it difficult to access the large PSRs. We note, however, that smaller and potentially more accessible PSRs occur on the intercrater plain (Fig. 1). The intercrater plain is a topographic depression spanning the nearside-farside transition and is composed mostly of Shoemaker and Shackleton ejecta with thinner intervening contributions from Faustini crater. Here, we present a new morphological map spanning 1600 km^2 between Shackleton and Shoemaker craters to highlight the units of the intercrater region and explore the area's potential for meeting Artemis objectives.

Methods: Morphological classification and boulder identification were developed using individual $\sim 50 \text{ km}$ -long, $0.5\text{m}/\text{pixel}$, LRO Narrow Angle Camera (NAC) image swaths and an averaged NAC mosaic distributed by MoonTrek [10]. Topographic analysis used LRO Lunar Orbiter Laser Altimeter (LOLA) digital elevation models (DEMs) at $\sim 5\text{m}/\text{pixel}$. Water-ice detection used Moon Mineralogy Mapper (M^3) data at $280\text{m}/\text{pixel}$ [4]. NAC image swaths were processed using the USGS Integrated Software for Imagers and Spectrometers (ISIS) 5.0.1. ArcGIS Pro 2.8.3 was used for visualization and mapping. Geospatial Data Abstraction Library (GDAL) was used for spatial calculation.

Morphology Evaluation: Morphological units are defined using geographic location, elevation, slope, NAC reflectance, and surface texture (Table 1). Average terrain roughness index (TRI) was calculated for each unit [11]. In this survey, we recognized four major morphological zones (Shackleton crater, Shoemaker crater, the intercrater plain, and ridges) with 21 morphological units (Fig. 2).

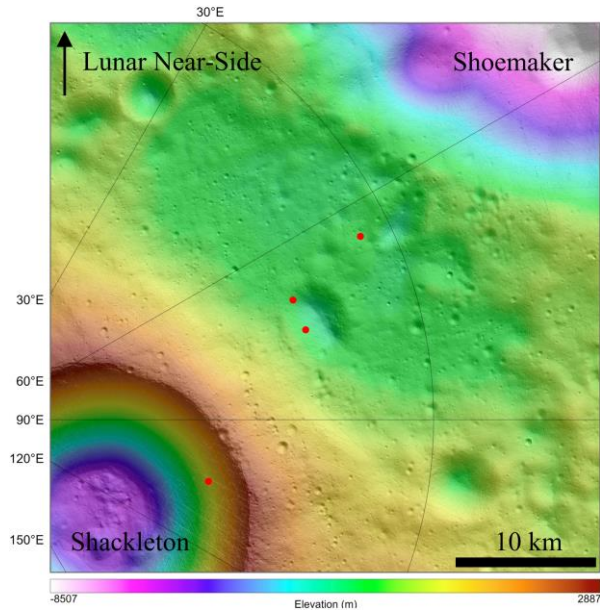


Figure 1: Topographic map of Shackleton, Shoemaker, and the intercrater region with overlain hillshade at 45° azimuth. Red points are M^3 water-ice detection [4].

Shackleton crater. Low-sloped circulate ridges along the rim of Shackleton contain high-reflectance blocky exposures that correlate to purest anorthosite (PAN) deposits [12]. Units are circumferentially distributed around Shackleton with lesser abundances of blocky features found outward and downslope from Shackleton's rim (units *ShRp* and *ShRtS*). Most of Shackleton's interior is permanently shadowed; however, portions of the upper crater walls are sunlit and composed of high-reflectance blocks that slump downward into the crater.

Shoemaker crater. The mapped area contains a portion of Shoemaker crater. Shoemaker's ejecta blanket, and regolith covering it, is divided into sloped units from the crater rim to the intercrater region.

Intercrater region. The intercrater region contains a diverse suite of rock exposures, boulders ejected from several locations, and underlying lineations of the lunar regolith. Contrasting reflectance units appear within areas of similar topography, possibly from newly exposed regolith or creep-like processes (units *PS* and *PmS*). Bright-reflectance rock exposures are observed near the contact with Shackleton and smaller primary craters (1-5 km diameter). Primary and secondary crater

walls and floors are mapped individually (units *CtW* and *CtF*), with smaller craters perched along primary and secondary crater walls (unit *PCtW*). Primary craters, secondary clusters, and perched craters (<1 km diameter) dominate the topography. Eroded crater walls or isolated steep slopes are defined as ridges. Smooth slopes with distinct reflectance from ridges are identified as plains.

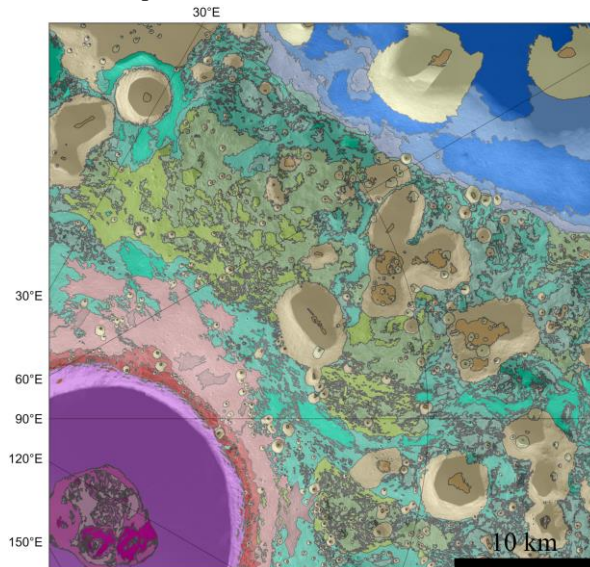


Figure 2: Geomorphic map of Shackleton crater, Shoemaker crater, and the intercrater region overlain on a NAC mosaic [10] and hillshade. Units are grouped by colour; Shackleton in red, Shoemaker in blue, intercrater plains in green, and cratered ridges in teal and yellow. Unit legend is described in Table 1.

Geomorphic Map: The most distinct geomorphic contacts are near the margin of Shackleton, where exposures of bright material appear mechanically transported from the slopes. Further distinction is found within the plains units, where trends of high-slope, high NAC reflectance, and *vice versa*, are found intermixed with slight differences in crater density (units *PmS* and *PS*). Ridges are orientated radially from Shackleton and Shoemaker near the unit zone margins and are generally less slumped with fewer effects of regolith creep toward Shoemaker. Craters 1-5 km diameter are positioned towards higher elevations, and provide areas of winter temperatures $\sim <110$ K. The most central primary craters (~ 5 km diameter, 600 m deep) within the plain are smoothly sloped ($<15^\circ$) to their floors and contain M^3 water-ice detection with rock exposures straddled along the rims. These 1-5 km diameter craters may also have excavated Shoemaker ejecta and massif material from beneath Shackleton ejecta.

Summary: Morphological units were mapped as a function of topography, NAC reflectance, and surface texture. The intercrater plains are mostly covered by

Table 1: Legend and classification of morphological units by slope, NAC reflectance, and TRI.

| Map Unit | Unit Description | Slope (Degrees) | NAC Reflectance | TRI |
|----------|---------------------------------|-----------------|-----------------|--------|
| CtW | Crater Wall | > 30 | 120 | 2.75 |
| PCtW | Perched Crater Wall | > 30 | 186 | > 10 |
| CtF | Crater Floor | ≤ 20 | 2 | 0.375 |
| PmS | Plain Moderately Hummocky | ≤ 20 | 99 | 1.125 |
| PS | Plain Hummocky | ≤ 10 | 23 | 0.125 |
| RhS | Ridge Highly Slumped | > 25 | 110 | 2 |
| RmS | Ridge Moderately Slumped | ≤ 25 | 56 | 0.75 |
| RS | Ridge Slumped | ≤ 10 | 35 | 0.125 |
| SohS | Shoemaker High Slope | > 30 | 160 | 3.2 |
| SotS | Shoemaker Transition Slope | ≤ 30 | 110 | 2 |
| SolS | Shoemaker Low Slope | ≤ 15 | 64 | 1.3 |
| ShRp | Shackleton Perched Rim | ≤ 15 | 96 | 3.1 |
| ShRhS | Shackleton Rim High Slope | > 30 | 166 | 4.75 |
| ShRtS | Shackleton Rim Transition Slope | ≤ 30 | 86 | 2.5 |
| ShRlS | Shackleton Rim Low Slope | ≤ 15 | 36 | 0.375 |
| ShW | Shackleton Wall | > 45 | 0 | > 10 |
| ShFhS | Shackleton Floor High Slope | > 30 | 0 | 4 |
| ShFmS | Shackleton Floor Moderate Slope | ≤ 30 | 0 | 1.25 |
| ShFlS | Shackleton Floor Low Slope | ≤ 15 | 0 | 0.125 |
| ShFPF | Shackleton Floor Perched Floor | ≤ 10 | 0 | 0.25 |
| ShacFF | Shackleton Floor | ≤ 10 | 0 | 0 |

Shackleton ejecta, but ridge exposures and ejecta from penetrating craters may have exposed Shoemaker ejecta and other lithologies from SPA impact melts or pre-Nectarian and Nectarian impacts [13]. Most blocky exposures are near the margins of Shackleton's units, which due to the proximity to Shackleton, may then contain lithologies of the original highland crust, the lunar magma ocean, and the cryptomare [13]. Units upon Shackleton contain blocky exposures, with some corresponding to previously identified PAN deposits. Units within the intercrater plain contain further blocky exposures and accessible areas ($<15^\circ$ slope) which may contain water-ice. Further investigation of these blocky exposures and of the overall accessibility within the intercrater region may provide a useful tool for Artemis exploration and for understanding the material and resource distribution on the lunar south pole [e.g., 14].

Acknowledgments: This work was supported by CLSE, LPI, USRA, and NASA-SSERVI to the 2021 Exploration Science Summer Intern Program.

References: [1] Spudis P. D. et al. (2008) *GRL*, 35, L14201. [2] Kring D. A. et al. (2021) *Adv. Space Res.* 68-11, 4691-4701. [3] Tye A. R. et al. (2015) *Icarus*, 255, 70-77. [4] Li et al. (2018) *PNAS*, 115-36, 8907-12 [5] Lemelin M. et al. (2021) *Planet. Sci. J.*, 2, 103. [6] NRC (2007) *The Sci. Context for Expl. of the Moon*, 107p [7] Anand M. et al. (2012) *Planet. Space Sci.*, 74, 42-48. [8] Mazarico E. et al. (2011) *Icarus*, 211, 114-120. [9] Speyerer E. J. and Robinson M. S. (2013) *Icarus*, 222, 122-136. [10] *USGS Lunar Map. and Model. Proj.* (2011) trek.nasa.gov. [11] Riley S. J. et al. (1999) *Interm. J. Sci.*, 5, 1-4. [12] Gawronska A. J. et al. (2020) *Adv. Space Res.* 66-6, 1247-1264. [13] Kring D. A. (2019) *NASA Expl. Sci. Forum.* [14] Bernhardt H. and Robinson M. S. (2021) *LPSC LII*, Abstract #1264