SULFATES ON MARS: A PYROCLASTIC AIRFALL MODEL FOR ORIGIN, EMPLACEMENT, AND INITIAL ALTERATION OF VALLES MARINERIS INTERIOR LAYERED DEPOSITS (ILD). James W. Head<sup>1</sup> & Lionel Wilson<sup>1,2</sup> <sup>1</sup>Dept. Earth, Environ. & Planet. Sci., Brown Univ., Providence RI 02912 USA. <sup>2</sup>Lancaster Environ. Centre, Lancaster Univ., Lancaster LA1 4YQ UK. (James Head@brown.edu) (l.wilson@lancaster.ac.uk)

it provides about: 1) mantle composition, source regions and differentiation processes, 2) magmatic eruptive processes, environments of sulfur exsolution and retention in effusive and explosive eruptions, sulfur speciation during eruptions and interaction with the atmosphere and aqueous and eolian transport, concentration and diagenesis, implications for climate evolution, 5) groundwater processes (solution, leaching from bedrock, transport, evaporative concentration and deposition, 6) postdepositional diagenesis and implications for the hydrological cycle, 7) host rock characteristics, sources, environments of emplacement, and 8) the role of sulfur in environmental conditions, liquid water acidity/alkalinity, and creation and maintenance of environments conducive or restrictive to formation and evolution of life.

shown that sulfur-related surface deposits are concentrated in time and space on Mars. Pioneering orbital remote sensing data [2] demonstrated that a period favoring nonacidic aqueous alteration and formation of phyllosilicates was followed by an acidic aqueous alteration environment in the Late Noachian-Early Hesperian which resulted in the emplacement of sulfate deposits, and was coincident with a period of significant volcanism. This acidic aqueous alteration environment was followed by the "atmospheric aqueous-free alteration" period of anhydrous ferric oxide formation during the Amazonian [2]. data [3] provided insight into the locations of the sulfaterich environments and the stratigraphic relationships between phyllosilicates and sulfates. Sulfates were concentrated in several distinctive environments (e.g., Meridiani-type and Valles Marineris-type layered deposits, intracrater clay-sulfates, polar gypsum deposits) and were often interbedded with phyllosilicates (particularly in Valles-type interior layered deposits, ILDs [3-6]) showing that the change in surface environments from nonacidic to acidic [3] may have been transitional.

What are the necessary requirements for the formation, evolution and preservation of sulfates? These conditions are highly specific due to high sulfate solubility and the environmental sensitivity of sulfates to phase transitions (temperature and humidity) [1] and include: 1) sources of S, 2) sources of liquid water and climates conducive to the derivation (leaching) and/or transport of S, 3) acidic environments resulting from S concentration in aqueous global warm and wet environmental conditions at the

Introduction: Sulfur is both ubiquitous and locally solutions, 4) sufficiently warm conditions to collect Sconcentrated on the surface of Mars and the martian sul- rich waters, and then to evaporate water and concentrate fur cycle [1] is of intense interest due to the information and deposit sulfates, 5) oscillating climate conditions to permit the interbedding of phyllosilicates and sulfates, 6) environments to explain their concentration in certain locations/settings, & 7) subsequent dry/cold climates necessary to preserve ancient sulfate deposits to the present.

Surface exploration of the Meridiani-type layered desurface, 3) atmospheric evolution and effects and con-posits by Opportunity showed that the Burns formation straints on climate, 4) surface sedimentary deposition, consisted of 18-25 wt% S, mixtures of Mg and Ca sulfates, and about 10% jarosite [7]. Together with sedimentary structures and stratigraphic relationships, these data were interpreted to indicate that it was emplaced in an eolian environment in which the sulfates formed in a groundwater upwelling low-pH aqueous evaporative environment, and underwent subsequent diagenesis. The process was envisioned to involve repeated upwellings of groundwater, periodically recharging local subaqueous depositional settings, and evaporating to concentrate the S and create the sulfate-rich deposits [7]. Pioneering the-The past several decades of Mars exploration have oretical and laboratory work [8-9] has shown that surface water and groundwater interactions with basalts can leach and concentrate S in solution, a primary prerequisite for further evaporative concentration and sulfate deposition. Modelling of the martian hydrological cycle [10] showed that an episodic evaporative environment at Meridiani could be formed by rainfall and infiltration in equatorial regions, lateral groundwater migration and leaching of sulfur during transport, and upwelling and evaporation at Meridiani, a general interpretation adopted to account for the Opportunity data [11].

Remote sensing characterization of the Valles Mari-Subsequent and complementary orbital remote sensing neris ILDs [3-6] showed that: 1) sulfate deposits were distinctly layered, 2) monohydrated sulfates were overlain by polyhydrated sulfates, 3) phyllosilicate and sulfate layers were interbedded, and 4) ILDs occur in kmsthick deposits exposed in eroded plateaus on the floors of Valles Marineris. While the origin and age of the ILDs are controversial [12], the formation of the sulfates is generally attributed to the same groundwater upwelling and evaporative deposition model as invoked at Meridiani [10]. Sulfur was concentrated by groundwater leaching from basaltic rocks, and further concentrated and deposited in sulfates by evaporative processes [3, 10].

While the groundwater concentration and evaporation scenario successfully accounts for a large number of requirements imposed by observations, it may not account for all documented environments of sulfate emplacement and preservation [3] on Mars. It also requires stringent sulfur mass-balance relations [12] and long-duration reconcile with climate models [13]. We explore the role ditions. 12) Sulfur speciation and atmospheric chemistry of magmatic processes [14] as an additional model for formation, evolution & preservation of sulfates on Mars.

Generation, Ascent & Eruption of Magma on Mars: As a basis for understanding the character and nature of volcanic eruptions in the martian environment (e.g., low gravity, currently low and historically evolving atmospheric pressure, presence of groundwater and ground ice) we have developed a theoretical and predictive framework for the generation, ascent and eruption of magma on Mars [14]. We have shown that basaltic plinian eruptions are highly favored (relative to Earth) [14-15] and explored the characteristics and dispersal of tephra [16] and gases from such eruptions in different locations and under different Patm conditions, and how this might lead to the observed distribution of fine-grained deposits [16]. The potential behavior and fate of S species during volcanic eruptions has also been investigated liquid water (co-precipitated and surface H<sub>2</sub>O melted by [17], including the role of H<sub>2</sub>SO<sub>4</sub> precipitates in surface H<sub>2</sub>SO<sub>4</sub>) and climates conducive to the derivation and/or melting & creation of aqueous acidic environments [18].

address the question: Can volcanic eruptions under marthe Valles Marineris Interior Layered Deposits? Obser-The Tharsis region is the most significant focus of volcanism on Mars, beginning in the Noachian and continuing through the Amazonian, with peak regional resurfacgeneral, and with increasing altitude (Tharsis) and decreasing Patm. 3) Any interaction of rising magma and groundwater/ground ice can contribute to explosivity [14]. 4) The martian environment favors the production of finer ash relative to Earth (~22 μm geometric mean for Mars), enhancing dispersal. 5) The presence and dominance of fine ash creates a profusion of readily available nucleation sites for condensation of co-erupted H2O and S species. 6) Airfall products are predicted to consist of tephra coated with condensed H2O and S species, producing extensive layered and graded deposits, and possible lapilli beds [15]. 7) The distribution of eruptive products is predicted to be focused latitudinally (equatorial in the case of Tharsis sources), with decreasing grainsize as a function of distance from the vent [16]. 8) Rapid formation of aerosols is predicted to occur, with global warming due to erupted S species being very short-lived, and global cooling ensuing until ambient conditions return [17]. 9) On the basis of calculated tephra dispersal and cooling times, temperatures of deposited tephra will decrease with distance from the vent, with the majority of deposit temperatures being close to that of the ambient atmosphere. 10) Magmatic exsolution of sulfur is favored by lower Patm and enhanced by higher altitude eruption sites (Tharsis). 11) Petrogenetic relations favor EPSL 294, 185.

time of formation, conditions that have been difficult to the co-exsolution of H<sub>2</sub>O and S species under these conpredictions favor H<sub>2</sub>SO<sub>4</sub> formation and widespread dispersal during and immediately following eruptions [18]. 13) Condensation and ensuing precipitation of H<sub>2</sub>SO<sub>4</sub> is predicted to melt any existing surface snow and ice, and to provide acidic aqueous surface environments favoring sulfate precipitation [18]. 14) Estimates of eruption duration and continuity readily predict deposit accumulations in excess of hundreds of meters to kilometers. 15) Oscillation of eruption conditions and S speciation may lead to alternate water-rich and sulfur-rich depositional conditions and interbedding of phyllosilicates and sulfates.

**Conclusions:** The characteristics of martian volcanism in the Tharsis region appear to meet the necessary requirements for the formation, evolution and preservation of sulfates in the ILD, including: 1) sources of sulfur (magmatic exsolution and precipitation), 2) sources of transport of sulfur (operates in a cold and icy climate; **Discussion:** On the basis of these considerations, we concentration by groundwater leaching of basaltic bedrock is not required), 3) acidic environments resulting tian conditions account for the characteristics of units in from sulfur concentration in aqueous solutions (provided by  $H_2SO_4$  precipitation and melting), 4) sufficiently vations consistent with such an hypothesis include: 1) warm conditions to collect S-rich waters (ponding of acidic meltwater created by H2SO4 precipitation), and then to evaporate water and concentrate and deposit sulfates (dry evaporative environments following eruptive ing in the Late Noachian-Early Hesperian [19]. 2) Explo- phases), 5) oscillating climate conditions to permit the sive plinian basaltic volcanism is favored on Mars in interbedding of phyllosilicates and sulfates (oscillatory conditions very plausible in the transition to lower  $P_{atm}$ ), 6) environments to explain their concentration in certain locations and settings on Mars (adjacent to long-lived Tharsis volcanic sources), and 7) subsequent dry and cold climatic conditions necessary to preserve the ancient sulfate deposits to the present (cold and icy conditions remain as  $P_{atm}$  evolves to present conditions).

> In order to test this hypothesis further, we are currently developing predictive eruption and tephra/volatile dispersal models that can serve as an interpretative framework to compare to the detailed characteristics and trends observed in the Valles Marineris ILDs [e.g. 3-6]. References: [1] King & McLennan (2010) Elements 6, 107; [2] Bibring et al. (2006) Science 312, 400; [3] Murchie et al. (2009) JGR 114, E00D06; [4] Roach et al. (2010) Icarus 206, 253; 207, 659; [5] Bishop et al. (2009) JGR 114, E00D09; [6] Wendt et al. (2011) Icarus 213, 86; [7] McLennan & Grotzinger (2008) The Martian Surface, Cambridge, 541; [8] Tosca & McLennan (2006) EPSL 241, 21; [9] Tosca et al. (2005) EPSL 240, 122; [10] Andrews-Hanna et al. (2007) Nature 446, 163; ibid (2010) JGR 115, E06002; Andrews-Hanna & Lewis (2011) JGR 116, E02007; [11] Arvidson et al. (2016) JGR 121, 1602; [12] Michalski & Niles (2012) Geology 40, 419; [13] Wordsworth et al. (2013) Icarus 222, 1; [14] Wilson & Head (1994) RG 32, 221; ibid, LPSC 51 1391; [15] Wilson & Head (2007) JVGR 163, 83; [16] Kerber et al. (2011) Icarus 216, 212; (2012) Icarus 219, 358; (2013) 223, 149; [17] Kerber et al. (2015) Icarus 261, 133; [18] Kreslavsky & Head (2020) LPSC 51 1828; [19] Carr & Head (2010)