

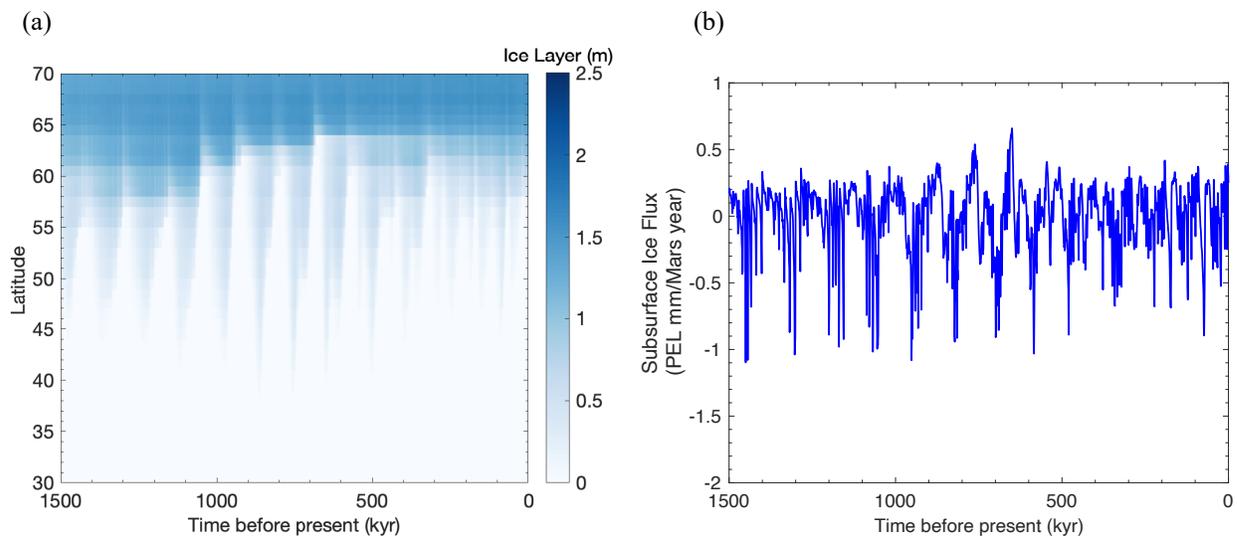
**Ice Reservoir Interactions Affect Polar Layered Deposit Evolution.** E. Vos<sup>1</sup>, O. Aharonson<sup>1,2</sup>, N. Schorghofer<sup>2</sup>, F. Forget<sup>3</sup>, E. Millour<sup>3</sup>, <sup>1</sup>Department of Earth and Planetary Sciences, Weizmann Institute of Science, Rehovot, Israel 76100; <sup>2</sup>Planetary Science Institute, Tucson, AZ 85719, USA, <sup>3</sup>LMD, Institut Pierre Simon Laplace Université Paris 6, France ([Eran.Vos@weizmann.ac.il](mailto:Eran.Vos@weizmann.ac.il))

**Introduction:** The climatic conditions which govern the formation of the North Polar Layered Deposits (NPLD) on Mars are influenced by interactions with vast mid-latitude ground ice deposits. Deciphering the paleoclimate record depends on understanding the behavior of the water sources and how they interact with each other and with the NPLD [1]. Both theoretical and observational work show that shallow subsurface ice is abundant in the mid-high latitudes, and exerts important influence on polar cap evolution. For instance, subsurface ice affects the seasonal CO<sub>2</sub> cover by storing heat during summer and releasing it at winter [2], due to the significantly increased thermal conductivity and thus inertia when ground-ice is present [3]. Using an innovative coupling a climate model and a long-term ice evolution model, in this work we explore how the ground-ice reservoir grows and recedes in response to changes in orbital elements and surface ice distribution in the past and how it, in turn, affects the NPLD accumulation and chemical composition.

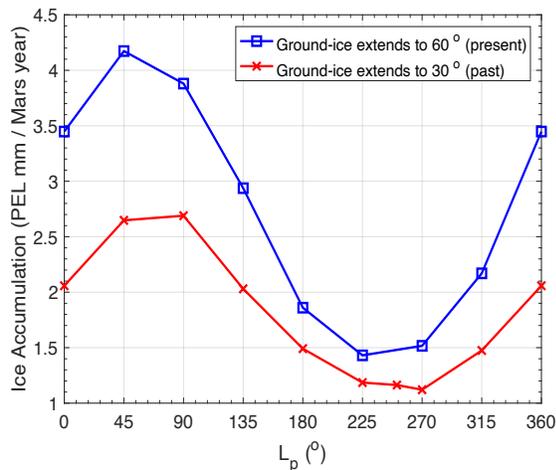
**Methods:** Our methods rely on creating a novel link between two state-of-the-art models that simulate aspects of the climate on two timescales— seasonal changes by a global climate model, and millennial changes by an ice evolution model.

1. *Global Climate Model (GCM):* we use the LMD-GCM to simulate and track the martian water cycle [4, 5]. Ice can precipitate in clouds or condense directly on the surface. The model includes the effects of the surface albedo of ice, resolves the vertical thermal structure in the atmosphere and subsurface down to ~30 m in depth, and models atmospheric dust. Surface ice initially resides in the tropics or in the pole, depending on the orbital configuration and the history. We allow the model run for 7 Mars years to reach steady state. We performed over 150 GCM simulations that span the orbital element variations and ice distributions considered, namely surface ice distributed as today as well as in the tropical regions as expected in the past.

2. *Mars Subsurface Ice Model (MSIM):* our 1D long-term ice evolution model relies on a previously published MSIM model [6]. It calculates the diffusive flux of vapor and ice deposited at the subsurface over thousands of years, based on near surface humidity and local properties. The relevant inputs are obtained from the GCM runs.



**Figure 1:** (a) Depth-integrated total ice layer thickness plotted over the last 1.5 million years as predicted by our long-term ice evolution model. Note both the obliquity and precession timescales control the ground ice growth and retreat. (b) Ice flux into the subsurface integrated over all latitudes. The flux is plotted in units of Polar Equivalent Layer (PEL), that is the equivalent depth of accumulation over an area the size of the current NPLD. For this simulation we assume 40% porosity.



**Figure 2:** Ice accumulation at the North Pole as a function of  $L_p$ , for today's orbital parameters. Note the  $\sim 50\%$  decrease in polar accumulation when ground ice is included in the model.

### Results:

Subsurface ice can affect the NPLD physical stratigraphy via several mechanisms. It acts as a water source, as well as a heat reservoir altering the energy budget [2]. The subsurface ice could also affect the polar chemical stratigraphy by changing the times and hence condensation temperature at which ice accumulates at the pole or by supplying water molecules with a different isotopic value [1]. We test and quantify these effects as follows. To test if ground-ice acts as a significant source we ran the ice evolution model for the last 3 Myr with changing orbital elements based on orbital integrations [7]. We track the evolution of the ice (Figure 1a), and integrate the subsurface flux over the latitudes that participate in the exchange (b). We find the magnitude of the flux to be of order mm's year, comparable to the polar accumulation rates computed in previous work [1, 8, 9].

To isolate the effect of the subsurface ice on the polar accumulation through changes in the thermal balance,

we constructed the GCM to have ice in the tropical region and ran simulations with the following orbital elements  $\epsilon = 25$ ,  $e = 0.093$  and a range of  $L_p$  values, where  $L_p$  is the solar longitude at perihelion. We ran two sets of simulations, one with ground-ice extends to  $60^\circ$  as present and one where the thermal inertia in the subsurface was set to mimic the presence of an ice sheet extending from  $30^\circ$  to  $80^\circ$  latitude in both hemispheres at a depth of 5 cm similar to past conditions. Figure 2 summarizes the results from both sets of simulations, and shows the presence of a massive ice sheet close to the surface reduces the flux by  $\sim 50\%$ . This effect is due to a change in the heat distribution preventing the high latitudes from becoming humid during the summer. Vos [9] showed that the temperature in the northern high latitude summer is an important determinant of NPLD accumulation when vapor is available from lower latitudes.

**Discussion:** Our results indicate that the subsurface ice affects the NPLD, acting both as significant vapor source, and altering the accumulation rates due to thermal effects (reduction of  $\sim 50\%$ ). As noted, the subsurface ice will further affect the chemical stratification of the polar deposits due to various processes [1].

### References:

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