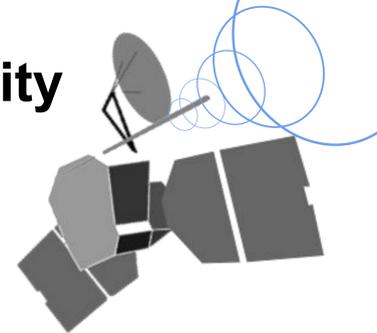


# Mars Subsurface Water Ice Mapping (SWIM): Radar Surface Reflectivity

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## What is SWIM?

The Subsurface Water Ice Mapping (SWIM) in the Northern Hemisphere of Mars Project supports an effort by NASA's Mars Exploration Program to determine in situ resource availability for future human exploration. Using several techniques and instruments, we are performing global reconnaissance mapping as well as focused multi-dataset mapping from 0° to 60°N. Results presented on this poster are focused on the radar surface reflectivity analysis of the SWIM project.

For results from the other datasets, see the other SWIM Project posters at this LPSC: **Bramson et al. (subsurface reflectors)** **Perry et al. (SWIM infrastructure)** **Hoover et al. (thermal analysis)** **Putzig et al. (geomorphology)** as well as the talk by Morgan et al. (overview) at 9:45 AM Friday morning

Final results will be presented at the next Human Landing Site Selection workshop. Our maps are being made available to the community on the SWIM Project website. Follow us on Twitter for project news and product release information.

<https://swim.psi.edu/> @RedPlanetSWIM

## Proposed Extension Activities

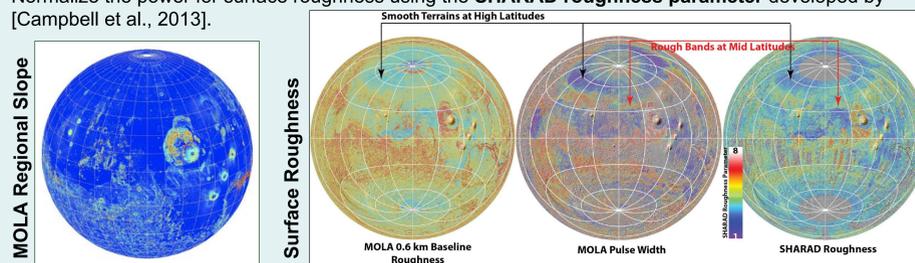
- Better calibrate the radar surface return data by accounting for spacecraft configuration using auxiliary information provided with SHARAD data
- Incorporate a new evaluation of SHARAD roughness
- Integrate results from MARSIS data

## Radar Surface Reflectivity

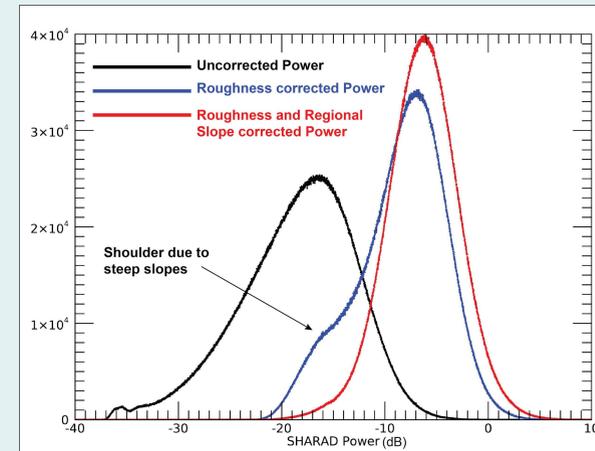
- SHARAD was designed to study the subsurface structure of Mars through the detection of reflections originating from boundaries between underground layers with contrasting dielectric properties.
- **HOWEVER**, Echoes returned from the surface of Mars also contain a wealth of useful information, including *surface roughness* and near-surface *Fresnel reflectivity*.
- Fresnel reflectivity is related to the *dielectric properties* and *density* of the shallow (< 5 m) subsurface.
- It is thus possible to isolate **the signature of ice in low power surface returns**, due to its **low density** [Mouginot, J. et al., 2010].

So how do we **isolate** the Fresnel reflectivity from the surface power return?

1. Limit ionosphere effects by excluding all daytime tracks.
2. Normalize the power for surface roughness using the **SHARAD roughness parameter** developed by [Campbell et al., 2013].



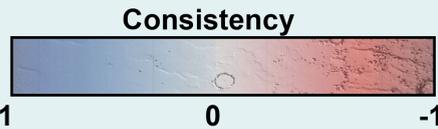
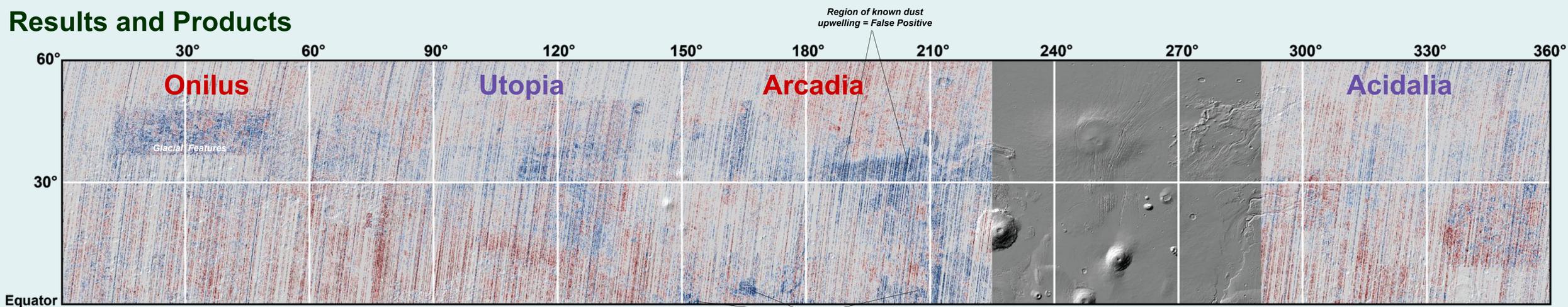
3. To further account for topographic effects at longer baselines, we correct for the loss of power due to regional slope using the median MOLA slope value over a Fresnel zone.
4. Finally, to account for additional MRO influences (spacecraft roll angle, solar panel configuration), we take the median value of corrected SHARAD returns sampled over 1/12° bins.



Above: Normalizing the SHARAD surface power for topography results in a narrowing of the power distribution and thus further constrains the influence of near-surface density.

Far Left: MOLA median regional slope values over a SHARAD fresnel zone. Left: Comparison of MOLA (Kreslavsky and Head, 2000: base-length and Neumann et al 2003: pulse width) versus SHARAD derived roughness datasets. All methods reveal similar surface features including those associated with periglacial landforms.

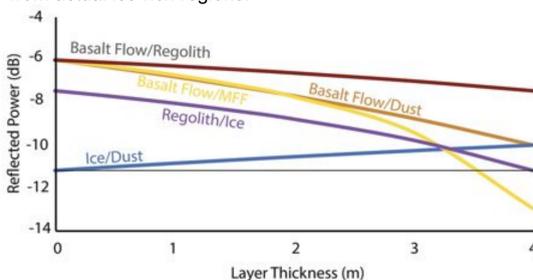
## Results and Products



**Martian Northern Hemisphere Radar Surface Reflectivity Ice Consistency Map.** As the surface power return is governed by the bulk density of the upper ~ 5m, we assume that low power (thus low density) values are **consistent with the presence of ice**. The consistency values displayed in the map are scaled according to the corrected power distribution presented up top. The power data show ice consistency in many distinct regions of the northern plains, including areas associated with the extensive glacial features present in 'Onilus'. The equator-straddling Medusae Fossae Formation (MFF) also exhibits high consistency values. Previous sounding studies of the MFF have revealed it to consist of icy or anonymously very low density (dry) material [Walters et al., 2007; Carter et al., 2008; Morgan et al., 2015].

## Measuring Near-surface Density

Due to the low density of ice relative to other, more typical (basalt-derived) Martian materials, high concentrations within the shallow subsurface will result in a distinctive low power return. However, like SHARAD dielectric measurements, ice mixed with progressively larger concentrations of other materials will result in non-unique reflectivity values. Through the integration of SHARAD surface power maps with other datasets, the SWIM approach provides a means to isolate 'false positives' (see center) from actual ice-rich regions.



Simple two-material model depicting the influence adding a secondary material (within the upper 4 m of the subsurface) has on the expected reflected radar power.

## Comparisons between Radar Reflectivity and Geomorphology

