

Evidence for recent wet-based crater glaciation in Tempe Terra, Mars?

Frances E.G. Butcher¹, M.R. Balme¹, C. Gallagher², N.S. Arnold³, S.J. Conway⁴, R.D. Storrar⁵, A. Hagermann¹, S.R. Lewis¹



¹The Open University, UK (frances.butcher@open.ac.uk), ²University College Dublin, Ireland, ³University of Cambridge, UK, ⁴CNRS, Laboratoire de Planétologie et Géodynamique, Nantes, France, ⁵Sheffield Hallam University, UK.



Evidence for basal melting of putative debris-covered glaciers in Mars' mid-latitudes is extremely rare.

- The glaciers are currently **frozen to their beds**, but has this always been the case?
- Eskers** (Fig 1) emerging from two mid-latitude glaciers [1-2] indicate at least two **localized melting events** beneath existing glaciers **~110-150 Myr ago** (Fig 2).

Eskers indicate past glacial melting.

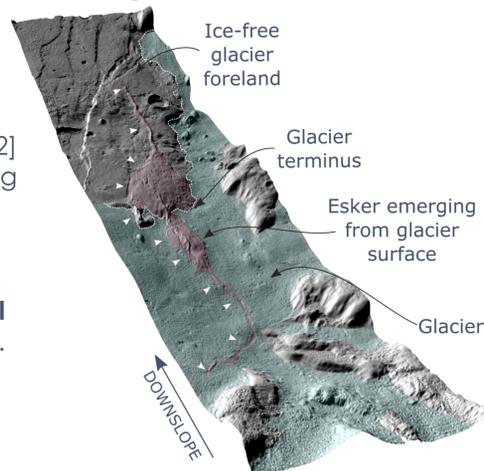
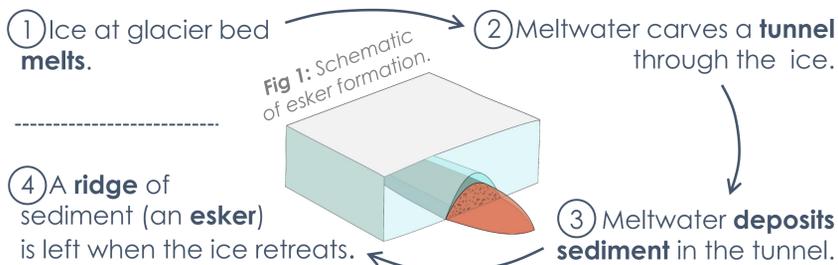
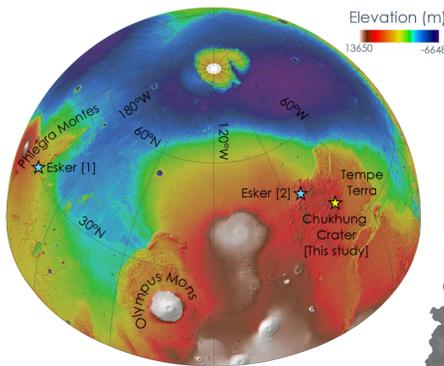


Fig 2: Esker (pink, 14 km long) emerging from a debris covered glacier (blue) in Tempe Terra [2]. Oblique shaded-relief map of HiRISE digital elevation model.

Are glacier-linked sinuous ridges in Chukhung Crater eskers?

- Chukhung Crater (38.47°N, 72.42°W) is a 45 km-diameter, Hesperian-aged [3] central pit crater [4] in Tempe Terra.



Unit interpretations

- Fim** Fresh impact material: (>200m craters).
- Hmu** Highland mantle unit: ice-rich airfall deposit?
- Vff** Viscous flow feature: remnant debris-covered glaciers.
- Gtr** Glacier-terminal ridges: glacial moraine ridges or pre-glacial crater wall slump deposits [e.g. 4].
- Rpu** Ridged plains unit: low-relief ground moraines?
- Ssr** Southern sinuous ridges: esker-like ridges extending from Gtr and Rpu.
- Tar** Transverse Aeolian Ridges occupying topographic lows: material possibly sourced from Isp.
- Ipu** Isolated pockmarked unit: patches of unknown origin; possible airfall deposit?
- Cpf** Central pit floor: closely-spaced aeolian bedforms: material sourced from Spm?
- Spm** Smooth plateau and mesas: alluvial or lacustrine deposits? Degrading into isolated mesas within Cpf.
- Usp** Upper smooth plains: layered plains continuous with inverted channel-like sinuous ridges; resistant fluvial deposits?
- Isp** Intermediate smooth plains dissected by broad sinuous valleys. Crater floor material or fluvial deposits, less resistant than Usp?
- Lsp** Lower smooth plains exposed within valleys dissecting Isp. Valley-filling deposits of fluvial or aeolian origin?
- Cwd** Crater wall deposits filling topographic lows within host crater wall & incised by anabranching valleys; fluvial deposits?
- Hcm** Host crater materials: crater rim, and central pit wall/rim elements.

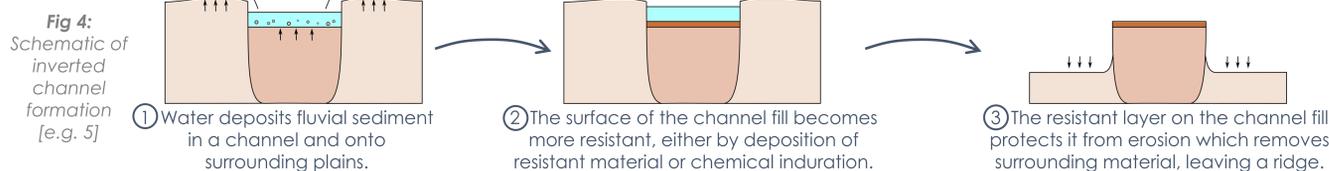
Structure

- Valley
- Impact crater rim

Fig 3: Geomorphic map of Chukhung Crater on CTX image basemap. Inset: MOLA elevation map of Mars' northern hemisphere showing the locations of Chukhung Crater & the two known glacier-linked eskers.

Chukhung Crater hosts two populations of sinuous ridges.

- Esker-like ridges** (Ssr, Fig 3) emerge from moraine-like deposits (Gtr & Rpu, Fig 3) bounding the termini of putative debris-covered glaciers (Vff, Fig 3) on the southern crater floor.
- Inverted channel-like ridges** (within Usp, Fig 3) extend from fluvial valleys on the northern crater wall. They formed **prior to glaciation** of the crater. Their formation **does not require glacial meltwater** (Fig 4).



The two sinuous ridge populations are morphologically distinct, supporting different origins.

- The esker-like ridges are **younger, more sinuous**, and have **sharper crests** than the inverted channel-like ridges (Fig 5).
- However, the ridges have **similar dimensions**, so differences in crest morphology could be due to **differences in degradation state** rather than formation mechanism.

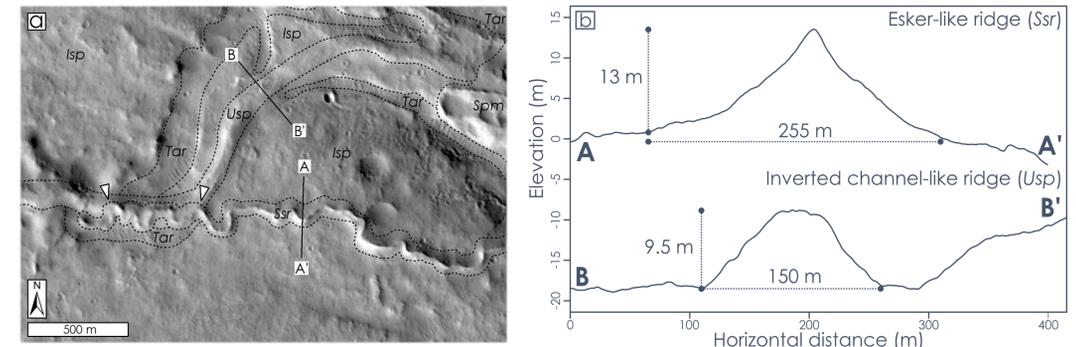


Fig 5: (a) Esker-like ridge (Ssr) superposing inverted channel-like ridge (Usp) (extent in Fig 3. CTX image P04_002577_2186_XN_38N072W), and (b) topographic profiles AA' and BB' from (a) extracted from digital elevation model generated from HiRISE images ESP_017477_2190 and ESP_018545_2190 [6].

The esker-like ridges ascend valley walls.

- Esker-forming meltwater can **ascend bed slopes** under hydraulic pressure in **subglacial tunnels** [8]. Ascent of valley walls (Fig 6b) is **inconsistent** with deposition under **gravity-driven flow** in subaerial fluvial channels.
- However, ascent of slopes could be **inherited from differential erosion** under the alternative inverted channel hypothesis, rather than a primary feature.

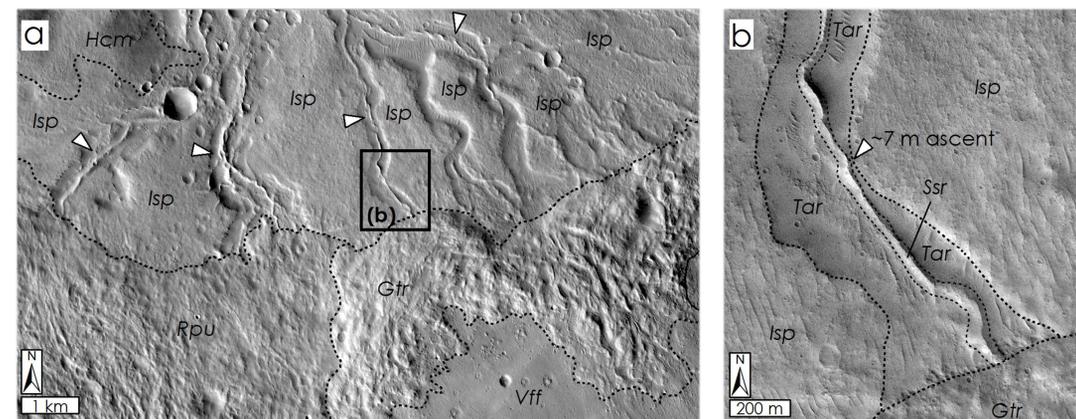


Fig 6: (a) Esker-like sinuous ridges (white arrows) emerging from moraine-like deposits (Gtr & Rpu) at glacier (Vff) termini. CTX image P04_002577_2186_XN_38N072W. (b) Esker-like ridge ascending a valley wall. HiRISE image ESP_023503_2185.

There are challenges for the esker hypothesis.

- The esker-like ridges could be a **second population of inverted channels**.
- Glacial deposits** (Vff, Gtr, Rpu) covering the southern crater floor hinder scrutiny of the relationship of the esker-like ridges to pre-glacial fluvial deposits.
- Eskers are **ice-contact deposits** but there is **no additional evidence** for past glaciation northward of the moraine-like deposits (Gtr & Rpu).
- There is one **esker-like ridge system on the northern floor**, where there is no evidence for glaciation.

Lessons from Chukhung Crater.

- Even where sinuous ridges emerge from **existing glaciers**, and where they have esker-like **non-slope-conforming topographic signatures**, conclusive identification as eskers is complicated by **similarities in form** between inverted channels and eskers [e.g. 8].
- Regional mapping** and **quantitative 3D morphometric analyses** [e.g. 2,9] should always be performed before an esker origin can be concluded. Such analyses are ongoing for Chukhung Crater.