

# When Moons Collide

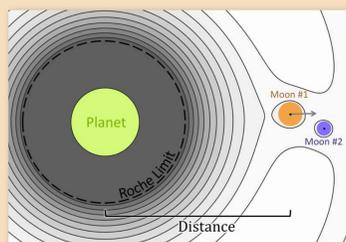
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## INTRODUCTION

Impacts between two orbiting satellites are a **natural consequence** of Moon formation [1,2,3]. Mergers between moonlets are especially interesting for the newly proposed multiple-impact hypothesis as these moonlets form from different debris disks and merge together to form the final Moon [3]. However, this process is relevant also for the canonical giant impact, as previous work shows that multiple moonlets can form from the same debris disk [1,4].

In this work we use Smoothed Particle Hydrodynamic code to estimate the **merger efficiency** of these impacts. The impacts occur within the planetary gravitational potential, therefore tidal forces alter the amount of mass that comprises the final moon. In order to directly compare our results with the standard impacts of two bodies in free space we perform reference simulations with no central potential.

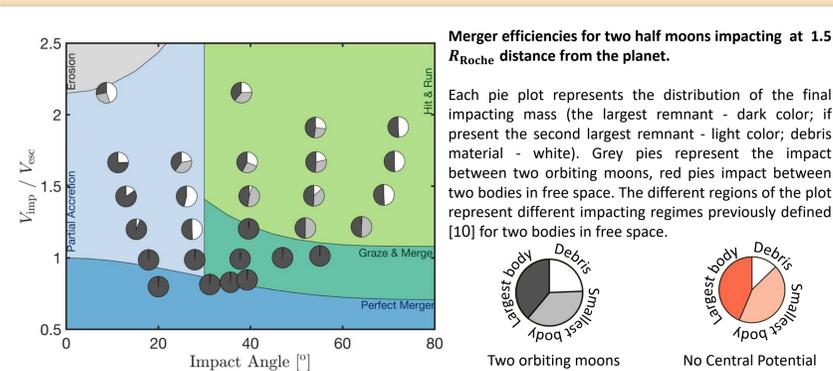


Previous simulations show that primordial debris disks and their accreted moonlets have different isotopic signatures, depending on the parameters of the collision with the planet [3]. Here we measure the **amount of mixing** between the two colliding moons in order to estimate the lunar heterogeneity [5].

The leading theory of the formation of the lunar anorthositic crust is the floatation of plagioclase minerals from a lunar magma ocean [8]. Moreover, gravity data from GRAIL reveal igneous intrusions that provide evidence for lunar radial expansion, consistent with a solidification of a 200-300 km-deep magma ocean [9]. Here we constrain the **amount of melting** in moonlet impacts in order to estimate the contribution to the lunar magma ocean.

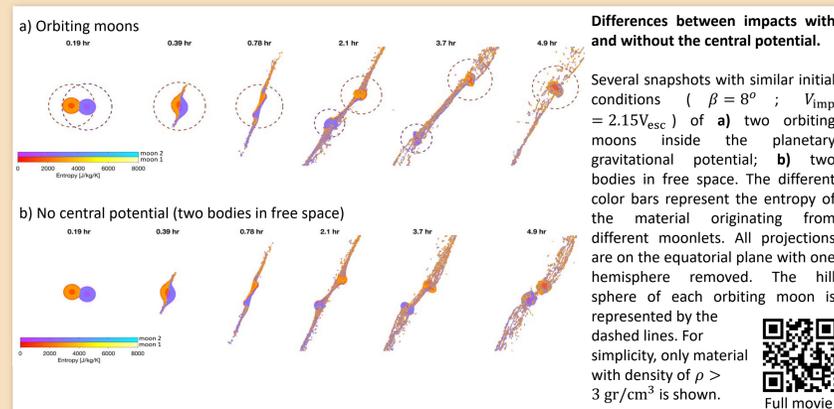
- How are moonlet collisions in a central planetary potential different?
- Do merged moonlets preserved heterogeneity?
- Are moonlet collisions consistent with the observed anorthositic crust?

## MASS EFFICIENCY

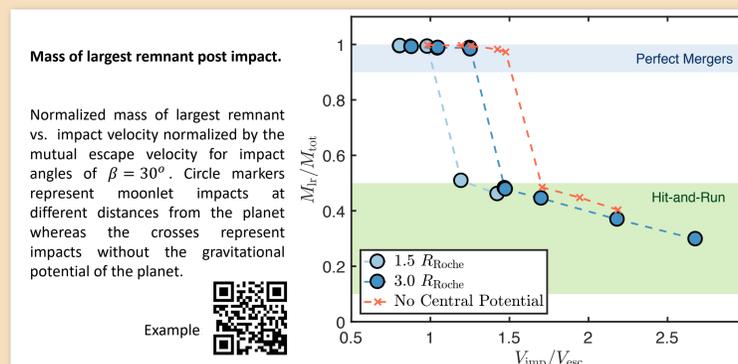


- Impacts in central potential promote moonlet loss.
- Moonlet erosion and debris generation are enhanced.

## HIT-AND-RUN

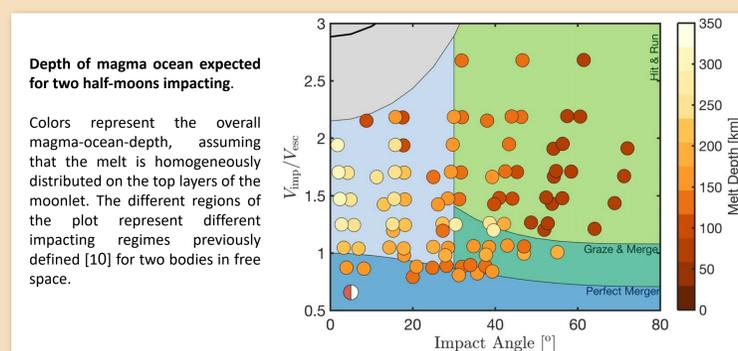


- Hit-and-run impacts are more abundant within a central potential.



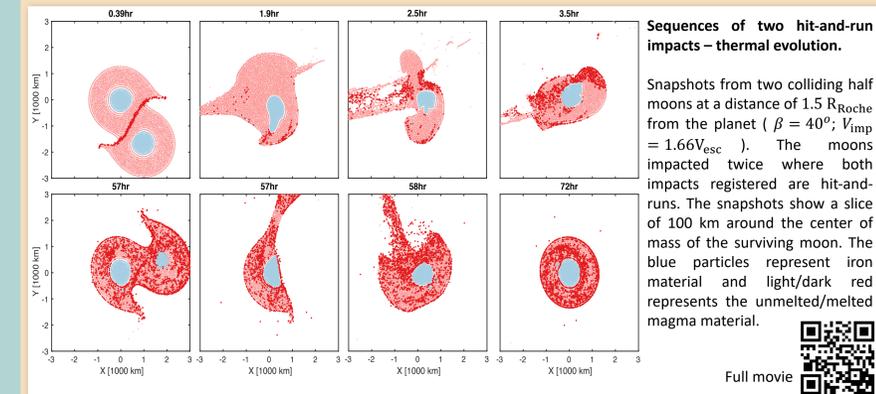
- Transition to Hit-and-run regime occurs at lower velocities closer to the planet (deeper in the potential well).
- Hit-and run impacts transfer small amount of angular momentum, the orbits are only modestly perturbed, therefore expect sequences of hit and runs.

## THERMAL STATE



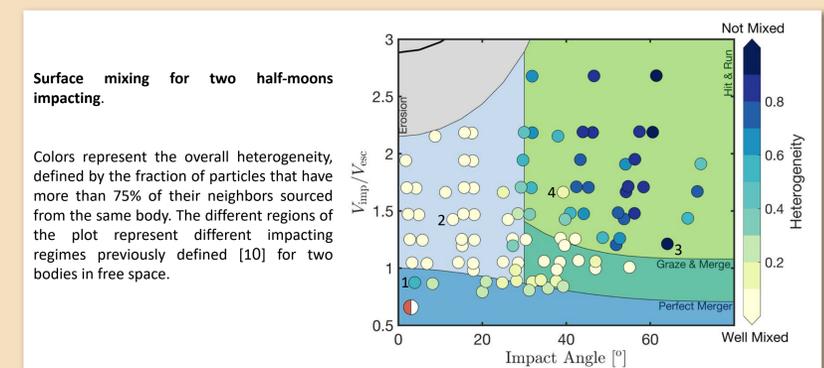
- Magma ocean (100's km) & anorthositic crust growth is expected

## THERMAL STATE cont.

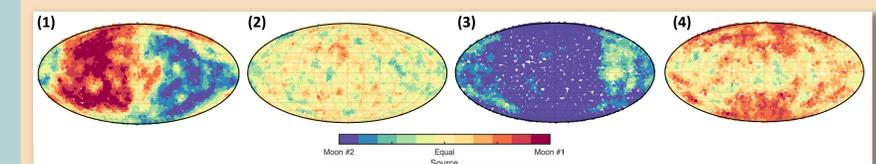


- Interaction of bodies of comparable size enables deeply sourced mantle to rise to the surface, decompress and melt.
- Sequences of impacts generate additional melt.

## SURFACE MIXING



- (1) Low velocity, head-on impact – not enough angular momentum for efficient mixing.
- (2) More typical accretory impacts are well mixed.
- (3) Hit-and-run impacts can create a heterogeneous surface.
- (4) However, sequences of hit-and-run impacts enhance mixing.



References:  
 [1] Ida S. et. al. (1997) *Nature*, 389, 353–357. [2] Jutzi M. and Asphaug E. (2011), *Nature*, 476, 69–72. [3] Rufu R. et. al. (2017), *Nature Geoscience*, 10, 89–94. [4] Salmon J. and Canup R. (2012), *Apl*, 760:83. [5] Robinson L. K. et. al. (2016), *Gca*, 199, 244–260. [6] Canup R. et. al. (1999), *Apl*, 117, 603–620. [7] Citron R. et. al. (2017), *DPS*, 508.05. [8] Elkins-Tanton L. T. et. al. (2011), *Epsl*, 304, 326–336. [9] Andrews-Hanna J. C. (2014), *Nature*, 514, 68–71. [10] Leinhardt Z. M. and Stewart S.T. (2012), *Apl*, 745:79.