

ASSESSMENT OF POTENTIAL REGIONAL VARIATIONS IN MODIFIED IMPACT CRATER MORPHOLOGY ON MARS. Robert A. Craddock¹, Lourenço Bandeira², and Alan D. Howard³, ¹Center for Earth and Planetary Studies, National Air and Space Museum, Smithsonian Institution, Washington, DC 20560 craddockb@si.edu, ²Centre for Natural Resources and the Environment, Instituto Superior Técnico. University of Lisbon, Av. Rovisco Pais 1049-001, Lisbon, Portugal, lpcbadeira@ist.utl.pt, ³Department of Environmental Sciences, PO Box 400123 Clark 205, University of Virginia, Charlottesville, VA 22904, <http://erode.evsc.virginia.edu>

Introduction: Modified impact craters represent some of the very first geologic features observed on Mars by spacecraft data [e.g., 1], yet we still have a poor understanding of how they were modified, and how these processes may have changed both temporally and spatially. Although arguments that they resulted from the effects of volatile-rich airfall deposits continue to persist [2] only erosional processes are capable of explaining the fact that different sized craters are preserved in different stages of modification. Diffusional processes, such as rainsplash or micrometeorite bombardment [3, 4], advective processes, such as surface runoff potentially associated with valley network development [3, 4], and aggradational processes, such as ejecta from the Hellas impact basin [5] have all been proposed. There are also some observations to suggest that the style of crater modification may have changed over time, too [4]. With modern spacecraft data it is possible to decipher the types of geologic processes that have occurred while quantifying the amount of modification that has taken place. We have pioneered the models and techniques for doing these types of quantitative analyses [3, 4, 6]. To date, however, we have only examined a small number of craters in limited areas. The geologic, climatic and hydrologic history of Mars is recorded in modified crater populations. Here we present the preliminary results to systematically extract and decipher that history.

Approach: Differences in crater morphology is reflected in the topography, which can be quantified using Mars Orbiter Laser Altimetry (MOLA) data. Although it would seem easier to extract topographic information from the released MOLA gridded products, using these products creates two problems. (1) Gridding averages the shot data within each grid cell and fills empty grid cells, reducing the accuracy. (2) The cross-track spacing between MOLA profiles is ~1 km near the equator, and some gaps are much larger. When such gaps occur near the edges of craters, under-sampling makes the gridded maps appear as though the crater was preferentially flattened by erosion, which would cause us to overestimate the amount of erosion that has taken place. As a result, we have developed our own semi-automated techniques to collect and quantify data from individual MOLA tracks for individual fresh and modified impact craters [6].

To begin our analysis, we accessed the global Mars crater database compiled by [7], which includes a qualitative assessment of the degradation state of individual impact craters ranging from a scale of 1 (“rimless”), 2 (“slightly elevated”), 3 (“some degradation/modification”), and 4 (“sharp”). Although subjective, we generally agreed with the classifications presented by [7] and applied it here. To check for differences in crater morphology between regions, we analyzed craters in the Sinus Sabaeus (MC-20, 315°-360°W, 0°-30°S), Iapygia (MC-21, 270°-315°W, 0°-30°S), Mare Tyrrhenum (MC-22, 225°-270°W, 0°-30°S), and Aeolis (MC-23, 180°-225°W, 0°-30°S) quadrangles. To look for potential differences due to latitude, we also analyzed craters in the Eridania (MC-29, 180°-240°W, 30°-65°S) quadrangle. The morphometric parameters presented in Table 1 include average values for the *crater depth*, which is the difference between the lowest elevation in the crater floor compared to the elevation of the surrounding crater exterior, *floor slope*, which is the calculated slope of the crater floor from the base of the wall to the center of the crater, average minimum curvature (*avg min curve*), which is the angle between the crater floor and the crater wall, the average maximum gradient (*avg max grad*), which is the average maximum slope of the crater wall, average maximum curvature (*avg max curv*), which is the angle between the crater wall and the surrounding exterior, average maximum elevation (*avg max ele*), which is the average maximum elevation of the landscape surrounding the crater out to one crater diameter, and the *rim height*, which is the height of the rim calculated from the average surrounding elevation and the total crater depth.

Results: In general, there is no obvious difference in morphometric parameters seen in modified impact craters located in different regions. For example, differences in the average maximum curvature has shown to be particularly diagnostic of craters that have been modified exclusively by linear diffusional creep versus craters that have also been modified by extensive surface runoff [3, 4, 6]. However, the minimum average value seen in Type 1 craters located in the Mare Tyrrhenum quadrangle is 10.9 ± 6.7 versus a high of 12.2 ± 7.2 for Type 1 craters located in the Iapygia quadrangle. While we cannot rule out the possibility that there are individual craters with unique morphologies that may be the result of differences in local lithol-

	DEPTH	FLOOR SLOPE	AVG. MIN CURV	AVG MAX GRAD	AVG MAX CURV	AVG MAX ELE	RIM HEIGHT
SINUS SABAEUS							
AVERAGE TYPE 1	424.4688	0.0289	9.8141	0.1267	11.8700	12.5177	77.3779
Standard Deviation Type 1	341.4358	0.0202	5.7502	0.0799	6.8009	7.3091	158.5398
AVERAGE TYPE 2	1096.1800	0.0545	14.6224	0.2472	17.5646	18.5680	172.0819
Standard Deviation Type 2	506.0065	0.0405	8.5864	0.0776	10.2421	10.9511	257.8756
AVERAGE TYPE 3	1254.0514	0.1230	7.9689	0.3613	9.7280	9.8548	162.9506
Standard Deviation Type 3	440.3839	0.0366	4.8502	0.0993	5.8358	6.2244	226.2438
AVERAGE TYPE 4	1158.8566	0.1476	6.0865	0.3698	7.5009	7.3612	147.6143
Standard Deviation Type 4	250.5244	0.0285	1.7479	0.1293	2.1959	2.2127	136.7709
IAPYGGIA							
AVERAGE TYPE 1	492.3537	0.0332	10.0381	0.1505	12.1562	12.8281	90.2259
Standard Deviation Type 1	367.4816	0.0225	6.0200	0.0840	7.2726	7.8167	189.5104
AVERAGE TYPE 2	1155.7900	0.0630	13.2958	0.2736	16.1134	16.9249	251.4145
Standard Deviation Type 2	538.9480	0.0367	8.3547	0.0844	10.0317	10.7596	310.2110
AVERAGE TYPE 3	1288.0029	0.1296	7.8410	0.3792	9.5506	9.6973	194.3229
Standard Deviation Type 3	434.5118	0.0392	4.5168	0.1145	5.4267	5.8512	257.3045
AVERAGE TYPE 4	1298.3502	0.1464	6.8147	0.3610	8.3753	8.2064	188.6785
Standard Deviation Type 4	414.7159	0.0186	1.8818	0.1338	2.2700	2.5137	133.5801
MARE TYRRHENUM							
AVERAGE TYPE 1	468.1848	0.0359	9.0601	0.1638	10.8881	11.4297	85.6204
Standard Deviation TYPE 1	331.3683	0.0228	5.5876	0.0837	6.6669	7.3749	166.2673
AVERAGE TYPE 2	974.7508	0.0500	13.0491	0.2606	15.6156	16.3349	233.1468
Standard Deviation TYPE 2	486.2858	0.0296	8.0941	0.0701	9.6861	10.2303	263.6800
AVERAGE TYPE 3	1271.9136	0.1284	8.0359	0.3684	9.7210	9.9282	206.9546
Standard Deviation TYPE 3	417.2479	0.0381	4.8633	0.1078	5.8286	6.2376	215.7420
AVERAGE TYPE 4	1139.7065	0.1372	5.9365	0.4220	7.2678	7.3173	178.1608
Standard Deviation TYPE 4	242.7316	0.0264	1.2333	0.0999	1.3824	1.5041	172.0588
AEOLIS							
AVERAGE TYPE 1	474.7582	0.0351	9.5400	0.1642	11.5176	12.0795	102.2380
Standard Deviation TYPE 1	327.4705	0.0222	6.0950	0.0829	7.3869	7.9364	193.6262
AVERAGE TYPE 2	919.6988	0.0471	12.2871	0.2645	14.7394	15.4780	235.6668
Standard Deviation TYPE 2	451.6106	0.0222	7.1917	0.0747	8.6093	9.3067	279.2618
AVERAGE TYPE 3	1248.5662	0.1285	7.6517	0.3729	9.2969	9.4273	186.9793
Standard Deviation TYPE 3	367.2849	0.0355	3.8268	0.0921	4.5421	4.8703	203.0818
AVERAGE TYPE 4	1164.3676	0.1162	6.6568	0.4085	7.9936	8.1557	148.0553
Standard Deviation TYPE 4	334.9862	0.0190	1.9714	0.0716	2.2346	2.5532	160.8445
ERIDANIA							
AVERAGE TYPE 1	376.1940	0.0267	9.3535	0.1251	11.4193	12.1320	65.6144
Standard Deviation Type 1	285.9383	0.0180	5.4197	0.0759	6.5178	7.0716	149.0631
AVERAGE TYPE 2	773.0100	0.0460	11.0959	0.2060	13.4550	14.4221	162.4578
Standard Deviation Type 2	431.0075	0.0265	7.0847	0.0657	8.4502	9.2418	233.9325
AVERAGE TYPE 3	1087.6500	0.0929	8.3945	0.3137	10.2346	10.4420	209.5433
Standard Deviation Type 3	492.8625	0.0346	4.6388	0.0936	5.5062	5.8893	235.0930
AVERAGE TYPE 4	1440.8470	0.1176	10.1665	0.3645	12.6917	13.3117	135.7143
Standard Deviation Type 4	532.0351	0.0554	7.9019	0.0861	9.8679	10.4795	131.3886

ogy or perhaps brief climatic episodes on Mars, our preliminary results indicate the the general morphology of the modified impact craters remains consistent at different stages of modification, even at higher latitudes. This may reflect the effects of an ancient climate that was global. A global climate may be the result of a primordial atmosphere that was dense enough to prohibit many regional or latitudinal variations in climate that are typical on the Earth. Essentially, modified impact craters may be the result of condensation and collapse of a primordial steam atmosphere and record the period of time before lakes, oceans, and valley network formation occurred. Con-

tinued statistical analyses of the morphometric parameters we derived along with a systematic analyses of the other modified impact craters can test this hypothesis.

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