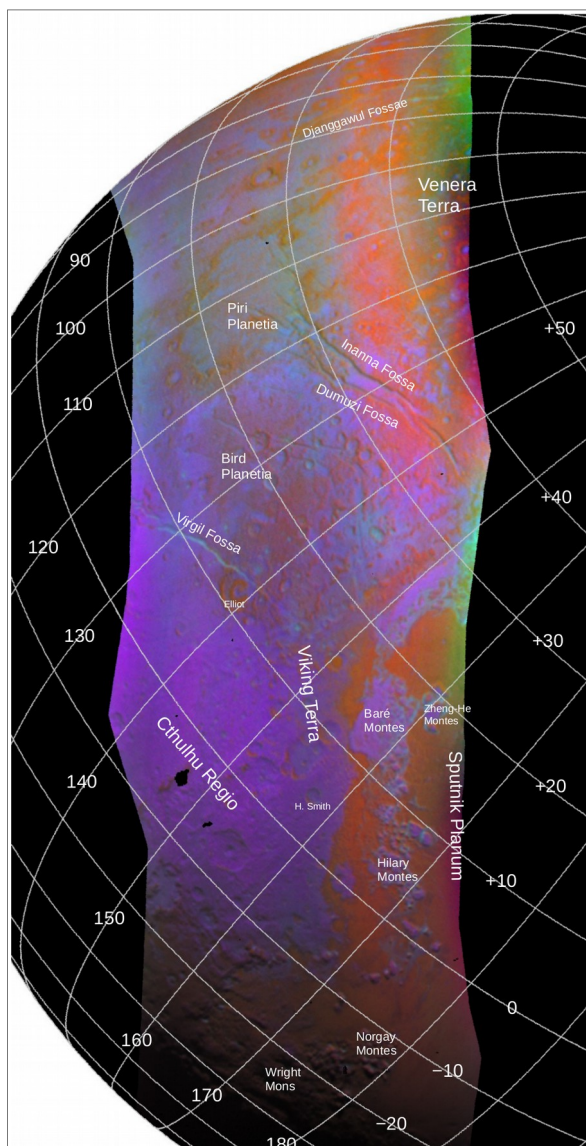


**HIGHEST SPATIAL RESOLUTION NEW HORIZONS LEISA SPECTRAL-IMAGING SCAN OF PLUTO.**

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**Observation:** The Linear Etalon Imaging Spectral Array (LEISA) component of New Horizons' Ralph instrument [1] provides spectral coverage from 1.25 to 2.5  $\mu\text{m}$ , at a resolving power ( $\lambda/\Delta\lambda$ ) of 240. LEISA's

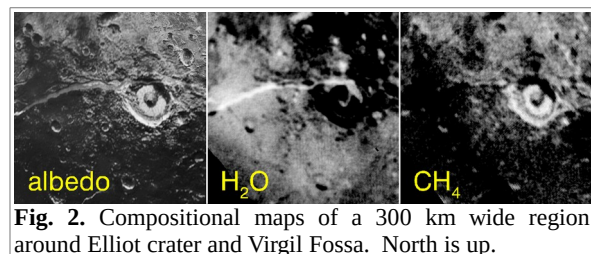


**Fig. 1.** False color image of the P\_LEISA\_HIRES scan constructed from the first few principal components computed from LEISA data. All feature names are informal.

spectral dispersion is provided by a wedge filter affixed to the detector so as to give each column of the array a unique wavelength. By scanning the array across the scene in the wavelength-dispersion direction while recording a series of images, each part of the scene is eventually imaged in each wavelength, enabling a multi-spectral image cube to be assembled.

The highest spatial resolution LEISA observation of Pluto, at about  $\sim 3$  km/pixel was obtained by New Horizons at around 10:56 UT 2015 July 14, from a range of 47,000 km. As shown in Fig. 1, the scan ran diagonally from the sunlit limb at mid-northern latitudes, southeast across the mountainous contact region between the large, dark, reddish expanse of Cthulhu Regio<sup>1</sup> and the extremely bright and volatile ice-rich Sputnik Planum. These features are so large and conspicuous that their existence had already been known from Earth-based observations, enabling the encounter sequence to be designed to reveal more about them.

**Results:** Maps of Pluto's various ices species were constructed from the LEISA observation by computing integrated areas, band depth ratios, or correlations with template spectra as described in [2]. These maps (e.g., Fig. 2) trace specific materials across Pluto's surface in a qualitative sense. More quantitative abundance maps await pixel-by-pixel analysis of the individual spectra, using radiative transfer models to account for multiple scattering and particle size effects, a computationally expensive process (see [3]).

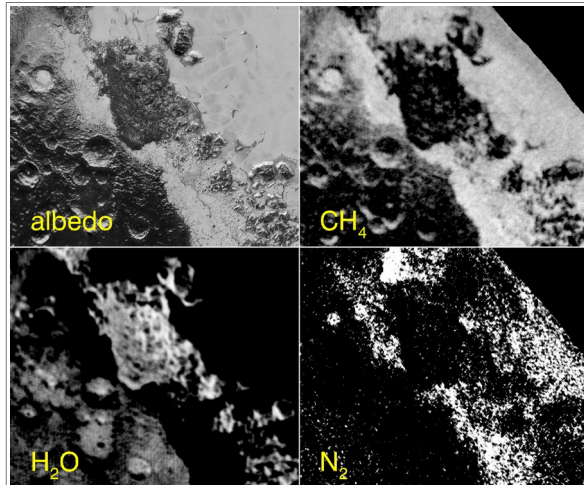


**Fig. 2.** Compositional maps of a 300 km wide region around Elliot crater and Virgil Fossae. North is up.

Fig. 2 shows  $\text{H}_2\text{O}$  and  $\text{CH}_4$  ice maps of the region around Elliot crater and Virgil Fossae, in comparison with an albedo map based on New Horizons LORRI

- 1 All feature names throughout this abstract and the associated talk are informal.

images [4,5,6]. The crater is strikingly depleted in  $H_2O$ , while the fossa is rich in that material, which is non-volatile at Pluto surface temperatures. The surrounding plains also show  $H_2O$  ice absorption but less so to the north of Elliot, where the  $H_2O$  gives way to  $CH_4$  ice. The flat floor of Elliot is very rich in  $CH_4$ , as are the north-facing portion of its inner rim and central peak.



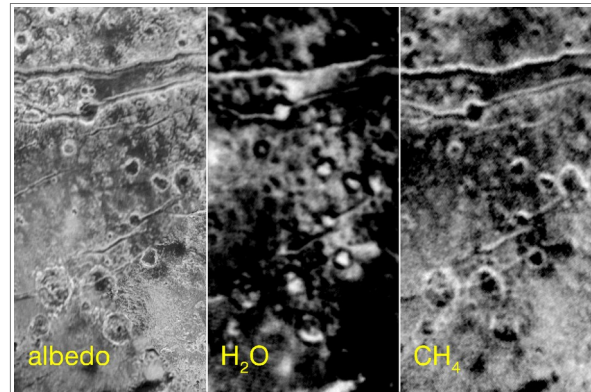
**Fig. 3.** Region extending from Viking Terra in eastern Cthulhu Regio at lower left to Sputnik Planum at upper right. The mountainous region at upper center is Baré Montes, with Zheng-He Montes near the top of the frame and Hilary Montes at lower right. Each pane is 350 km tall and north is up.

Another interesting region appears in Fig. 3, with maps of albedo,  $H_2O$ ,  $CH_4$ , and highly volatile  $N_2$  ice. In dark Cthulhu Regio at lower left,  $H_2O$  ice has a patchy distribution, and there is little or no  $N_2$  ice.  $CH_4$  occurs there, but mostly just on north-facing slopes. No  $H_2O$  is seen in Sputnik Planum, but that area is rich in both  $N_2$  and  $CH_4$ . Interestingly, the  $N_2$  ice absorption in Sputnik does not appear to be spatially uniform (unfortunately this is a noisier map owing to the weakness of the  $N_2$  absorption band at  $2.15 \mu m$ ). A region just north of Baré Montes shows particularly strong  $N_2$  absorption. An unnamed crater with a bright floor stands out at upper left. The bright floor material looks similar to Sputnik Planum, rich in both  $CH_4$  and  $N_2$ .

Despite its low visible wavelength albedo, Baré Montes shows reasonably strong  $H_2O$  ice absorption and little  $CH_4$  or  $N_2$ .  $H_2O$  ice was expected in mountainous regions, since the more volatile ices are too ductile for large, enduring edifices to be built of them. But its distribution in the various mountain ranges in this region is curious, tending to be associated with south-facing slopes. It appears to be masked by  $CH_4$  ice on north-facing slopes.

At mid-northern latitudes, as in Fig. 4, the distributions of  $CH_4$  and  $H_2O$  ices often appear anti-correlated.

Many craters show bright rims with strong  $CH_4$  absorption, while their dark floors show  $H_2O$  absorption. A possible explanation for this  $CH_4$  distribution is that it preferentially condenses from the atmosphere on topographic highs and regions with high albedos.



**Fig. 4.** Region extending from Bird Planitia north to Dumuzi and Inanna Fossae. Each pane is 400 km tall and north is up.

This co-occurrence of  $H_2O$  absorption with visually dark regions is seen widely across Pluto, with Baré Montes and Cthulhu Regio providing good examples in Fig. 3. It is unclear whether the  $H_2O$  plus tholins responsible for darkening the crater floors actually accumulate there (perhaps arriving as wind-blown dust), or are generally representative of Pluto's subsurface materials but are seen to outcrop only in regions such as crater floors where they are least likely to become coated by volatile ices, owing to the local thermal environment.

**References:** [1] Reuter et al. (2008) SSR, 140, 129-154. [2] Grundy et al. (2016) Science, submitted. [3] Protopapa et al. (2016) this conference. [4] Cheng et al. (2008) SSR, 140, 189-215. [5] Stern et al. (2015) Science, 350, 292. [6] Moore et al. (2016) Science, submitted.