

HIGH-PRECISION RUTHENIUM ISOTOPE MEASUREMENTS FOR CONSTRAINING LATE ACCRETION. M. Fischer-Gödde¹, C. Burkhardt² and T. Kleine¹. ¹Institut für Planetologie, Westfälische Wilhelms-Universität Münster, Wilhelm-Klemm-Str. 10, 48149 Münster, Germany. Correspondence: m.fischer-goedde@uni-muenster.de, ²Origins Laboratory, Department of the Geophysical Sciences and Enrico Fermi Institute, The University of Chicago, IL 60637, USA.

Introduction: Excess abundances of highly siderophile elements (Re, Os, Ir, Ru, Pt, Rh, Pd, Au) observed in the Earth's mantle are commonly explained by the addition of a late veneer of primitive meteoritic material to the mantle after core formation was complete [1-3]. The composition of the late veneer has been inferred from the comparison of Os isotope compositions and relative HSE abundances in Earth's primitive mantle to those of chondritic meteorites [4-11]. However, while the ¹⁸⁷Os/¹⁸⁸Os ratio of the Earth's primitive mantle overlaps with those of ordinary and enstatite chondrites, relative HSE abundances are inconsistent with a late veneer this composition. In particular the suprachondritic Ru/Ir of the primitive mantle cannot be accounted for by late accretion of known chondrites [e.g. 4,8].

The composition of the late veneer may also be assessed through the analysis of lunar impact melt rocks. These rocks bear testimony to the composition of bodies impacting the Earth and Moon since their formation about 4.4 Ga ago. Relative HSE abundances and ¹⁸⁷Os/¹⁸⁸Os ratios of some lunar impact rocks reveal the presence of both chondritic and differentiated meteoritic impactor material [12-15]. On this basis it has been proposed that the HSE composition of the Earth's primitive mantle has been established by late accretion of predominantly carbonaceous chondrite material, mixed with a minor portion of differentiated iron meteorite-like material [15]. The fractionated iron meteorite-like component can account for the elevated Re/Os (and hence the observed ¹⁸⁷Os/¹⁸⁸Os), Ru/Ir and Pd/Ir of the Earth's mantle in comparison to carbonaceous chondrites. More recently, the relative abundances of Se, Te and S have also been used to link the late veneer to a specific type of chondrite [16]. The chalcogen systematics suggest that the late veneer predominantly consists of carbonaceous chondrite-like material, consistent with the composition inferred from the lunar impact melt rocks.

The Ru isotope compositions of meteorites and the Earth's mantle provide new constraints on the composition and origin of the late veneer. Because almost all meteorites are characterized by nucleosynthetic Ru isotope anomalies, caused by a deficit in *s*-process nuclides relative to terrestrial Ru [17,18], Ru isotopes are a sensitive tracer for the composition of the late veneer and to assess whether known meteorites can be

the source of the late veneer. This is possible because almost the entire Ru in the Earth's mantle derives from the late veneer, so that only meteorites (or a combination thereof) having a terrestrial Ru isotope composition can be source of the late veneer.

Here we use high-precision Ru isotope measurements to constrain the origin of the late veneer by investigating a comprehensive set of meteorites with a special focus on those meteorite groups that are potential sources of the late veneer.

Analytical techniques: Iron meteorite samples (~0.05-0.5g) were digested in reverse aqua regia in Saville beakers on a hot plate. Analytical techniques for Ru separation and purification are provided in [20]. Chondrite powders (0.4-0.6 g) were digested in reverse aqua regia inside Carius tubes at 230 °C for 48 hours and purified using the same techniques as for the iron meteorites. Three Allende samples were dissolved using an alkaline fusion technique as described in [21]. Ruthenium from fluxed samples was purified by distillation in a PFA distillation unit.

Ruthenium isotopes were measured on a ThermoScientific Neptune *Plus* MC-ICPMS at the University of Münster. The measurements were typically performed with total ion beam intensities of $\sim 7.5 \times 10^{-11}$ – $\sim 4.4 \times 10^{-10}$ A, obtained for a 100 ppb Ru standard solution. Isobaric interferences of Mo and Pd on Ru masses 96, 98, 100, 102 and 104 were corrected by monitoring ⁹⁷Mo and ¹⁰⁵Pd. Measured Ru isotope ratios were normalized to ⁹⁹Ru/¹⁰¹Ru using the exponential law and are reported in $\epsilon^{100}\text{Ru}$ units representing the deviation in parts per 10⁴ from the terrestrial Ru isotope composition. Accuracy and precision of the Ru isotope measurements were evaluated by replicate digestions and multiple analyses of the NIST 129c steel doped with ~500 to ~2000 ng Ru from an Alfa Aesar standard solution. The external reproducibility (2 s.d.) of the Ru isotope analyses obtained for 49 individual measurements from 9 different digestions of NIST 129c over a period of the last two years is $\pm 0.45 \epsilon^{96}\text{Ru}$, $\pm 0.51 \epsilon^{98}\text{Ru}$, $\pm 0.13 \epsilon^{100}\text{Ru}$, $\pm 0.14 \epsilon^{102}\text{Ru}$, and $\pm 0.35 \epsilon^{104}\text{Ru}$.

Results: Most meteorites and chondrites are characterized by well-resolved negative $\epsilon^{100}\text{Ru}$ anomalies (Fig. 1). The IAB iron meteorites and EL enstatite chondrites are the only meteorites not showing clearly resolved Ru isotope anomalies, although the EL chondrites are probably also characterized by a small ¹⁰⁰Ru

deficit. We are currently analyzing additional enstatite chondrites to assess whether their Ru isotope composition is different from that of the Earth's mantle.

A prerequisite to utilize Ru isotope measurements for constraining the origin of the late veneer is that the true bulk Ru isotope composition of primitive chondrites is measured. However, many of these contain presolar grains, which are not easily dissolved using standard digestion techniques. We have found that Allende samples digested in reverse aqua regia in Carius tubes display larger and more variable $\epsilon^{100}\text{Ru}$ anomalies in comparison to Allende samples digested by alkaline fusion (Fig. 2). Obviously, Carius tube digestion is insufficient to dissolve all presolar components present in primitive chondrites, highlighting the importance of using a total digestion technique for Ru isotope studies. All chondrites investigated in the present study were either digested by alkaline fusion or are equilibrated samples devoid of presolar grains.

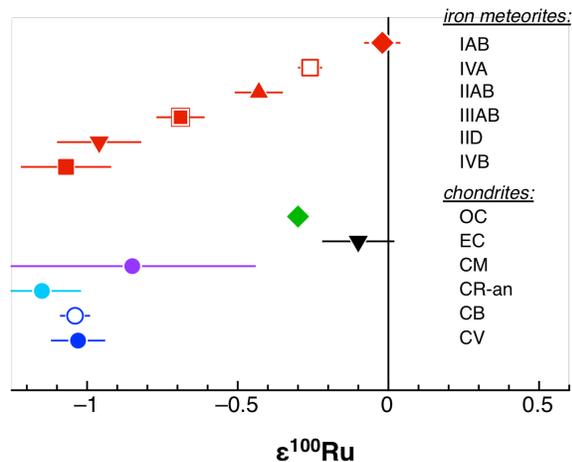


Fig. 1: $\epsilon^{100}\text{Ru}$ data for iron meteorites and chondrites.

Discussion: All meteorites investigated in the present study (except the IAB irons and possibly the enstatite chondrites) exhibit nucleosynthetic Ru isotope anomalies and, therefore, cannot be the source of the late veneer. Moreover, because all meteorites are characterized by a deficit in *s*-process Ru relative to the Earth's mantle, a combination of different meteorites also does not result in the terrestrial Ru isotope composition. Therefore, neither the investigated carbonaceous chondrites nor a combination of carbonaceous chondrites with any of the investigated magmatic iron meteorites can represent the material added to the Earth's mantle as a late veneer. Based on the Ru isotope results only the IAB irons and perhaps the EL chondrites are possible sources of the late veneer, but their chemical compositions do not match the HSE and chalcogen systematics of the Earth's mantle.

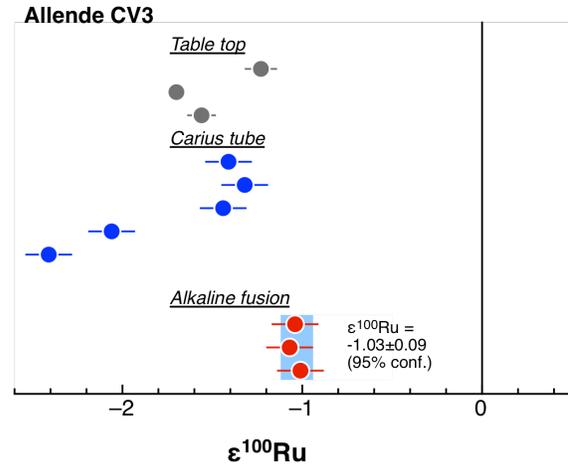


Fig. 2: $\epsilon^{100}\text{Ru}$ data for replicate analyses of Allende using different digestion methods. Data for Carius tube digestions (blue circles, this study) and digestions in teflon vials on a hot plate (grey circles, [17]) display more scatter and larger isotope anomalies compared to data obtained for fused samples (red circles).

Thus, the Ru isotope results, if substantiated by additional analyses of a more comprehensive suite of primitive carbonaceous and enstatite chondrites, suggest that the late veneer derives from a different population of bodies than known meteorites. One possibility is that the late veneer consists of the material remaining in the terrestrial planet region after the last giant impact occurred on Earth. This would be consistent with some recent dynamical models, in which volatile-rich, carbonaceous chondrite-like material is transported into the inner solar system during the final stages of terrestrial planet accretion [22].

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