

NEUTRON DOSIMETRY AND NEUTRON CAPTURE MODEL FOR PALLADIUM-SILVER CHRONOMETRY OF IRON METEORITES. M. Matthes¹, M. Fischer-Gödde¹, T. S. Kruijer¹, I. Leya² and T. Kleine¹, ¹Institut für Planetologie, Westfälische Wilhelms-Universität Münster, 48149 Münster, Germany, ²Space Research and Planetology, University of Bern, Bern, Switzerland (max.matthes@uni-muenster.de).

Introduction: The short-lived ^{107}Pd - ^{107}Ag system ($t_{1/2} = 6.5$ Ma) is a powerful chronometer for constraining the accretion and cooling histories of iron meteorite parent bodies [1-4]. However, exposure of the iron meteoroids to galactic cosmic rays (GCR) may have modified the Ag isotope compositions. This not only involves neutron capture-induced burnout of Ag isotopes, but also production of Ag via neutron capture on Pd isotopes [5]. Therefore, the net GCR effect on Ag isotope compositions increases with increasing Pd/Ag. Hence, although high-Pd/Ag iron meteorites can be very well dated using the Pd-Ag system, these meteorites may also show the largest GCR-induced shifts on $^{107}\text{Ag}/^{109}\text{Ag}$. Thus, applying the Pd-Ag system to iron meteorites requires a quantitative assessment of neutron capture effects in each sample.

The objectives of this study are to evaluate the significance of GCR-induced shifts on $^{107}\text{Ag}/^{109}\text{Ag}$ ratios in iron meteorites and to develop a method for correcting these effects. Such a correction method requires an independent neutron-dose proxy, and a model linking this dosimeter to the effects on $^{107}\text{Ag}/^{109}\text{Ag}$. To this end we have extended the model calculations of GCR-induced effects on Pd-Ag systematics from Leya and Masarik [5] with calculations of neutron capture effects on Pt isotopes, which provide a powerful neutron-dose proxy for iron meteorites [6,7]. The updated model makes it possible to correct GCR-induced effects on $^{107}\text{Ag}/^{109}\text{Ag}$ using Pt isotope compositions. To test the model, we obtained combined Pd-Ag and Pt isotope data for several splits of the iron meteorites Ainsworth (IIAB), Carbo and Rodeo (IID) as well as Grant (IIIAB). The first two of these samples are among the most strongly irradiated iron meteorites and, hence, ideally suited to study GCR-induced shifts on $^{107}\text{Ag}/^{109}\text{Ag}$ ratios, and for testing the newly developed correction method.

Analytical methods: All samples were polished using SiC abrasives, cleaned by ultrasonication in deionized water, and leached in warm 6 M HCl. After dissolution of the samples (1-3 g) in hot reverse aqua regia, three different aliquots were taken for the determination of Pd and Ag concentrations and Pt isotope compositions. A three-column ion exchange procedure, adapted and slightly modified from [3,8], was used for the purification of Ag. The separation of Pt followed the method of [6]. All isotope measurements were performed on the Thermo Scientific[®] Neptune Plus MC-

ICPMS in the Institut für Planetologie, using either an ESI APEX-Q (for Ag) or a Cetac Aridus II desolvating system (for Pt). Prior to measurement the Ag fractions were doped with Pd for mass bias correction relative to $^{108}\text{Pd}/^{106}\text{Pd} = 0.97237$. The Ag isotope data for samples are reported in $\epsilon^{107}\text{Ag}$ as the parts per 10,000 deviations from the mean $^{107}\text{Ag}/^{109}\text{Ag}$ obtained for measurements of bracketing runs of the NIST 978a standard. The reproducibility of the Ag isotope measurement is $\pm 2 \epsilon^{107}\text{Ag}$ (2s.d.), as estimated from repeated measurements of the NIST 129c steel. The Pt isotope measurements followed the protocols described in [6].

Neutron capture model: The model calculations follow those presented earlier [5]. We consider thermal and epithermal neutron capture reactions as well as reactions induced by fast particles (i.e., protons and neutrons with energies of a few MeV) on Ag, Pd, and Cd to fully cover all possible reaction pathways. The modeled GCR-induced shift on $\epsilon^{107}\text{Ag}$ depends on the Pd/Ag ratio, because neutron capture on ^{106}Pd and ^{108}Pd produces, after subsequent β^- -decay, ^{107}Ag and ^{109}Ag . In addition, the GCR-induced effects also depend on the neutron dose, which itself is a function of exposure time, pre-atmospheric radius of the meteoroid, and the shielding depth of the sample in the pre-atmospheric object. Consequently, for correcting GCR-induced effects on $\epsilon^{107}\text{Ag}$ two parameters must be considered, namely the Pd/Ag ratio and the neutron dose. For the latter we use the isotopic shifts in Pt isotopes, which are well suited for this task [6,7]. The model gives, for a given Pd/Ag, a linear regression of the type: $\epsilon^{107}\text{Ag}_{\text{GCR}} = \text{Offset} + \text{Slope} \times \epsilon^{196}\text{Pt}$, where both *Offset* and *Slope* are functions of the Pd/Ag ratio. Thus, the GCR-induced shift on $\epsilon^{107}\text{Ag}$ can be calculated for each sample using its measured Pd/Ag and $\epsilon^{196}\text{Pt}$.

Results: With the exception of Rodeo and two Grant specimens, all samples exhibit well-resolved Pt isotope anomalies. These indicate a variable neutron dose in the investigated samples, with Ainsworth having the highest neutron dose, followed by Carbo and Grant. In spite of high $^{108}\text{Pd}/^{109}\text{Ag}$ in the investigated irons, only Rodeo and Grant show ^{107}Ag excesses; Ainsworth and Carbo display negative $\epsilon^{107}\text{Ag}$ values. This is surprising, because the high $^{108}\text{Pd}/^{109}\text{Ag}$ of the irons should have led to ^{107}Ag excesses through the decay of live ^{107}Pd . Moreover, no linear correlation is obtained between $\epsilon^{107}\text{Ag}$ and $^{108}\text{Pd}/^{109}\text{Ag}$ for any of the investigated samples (Fig. 1).

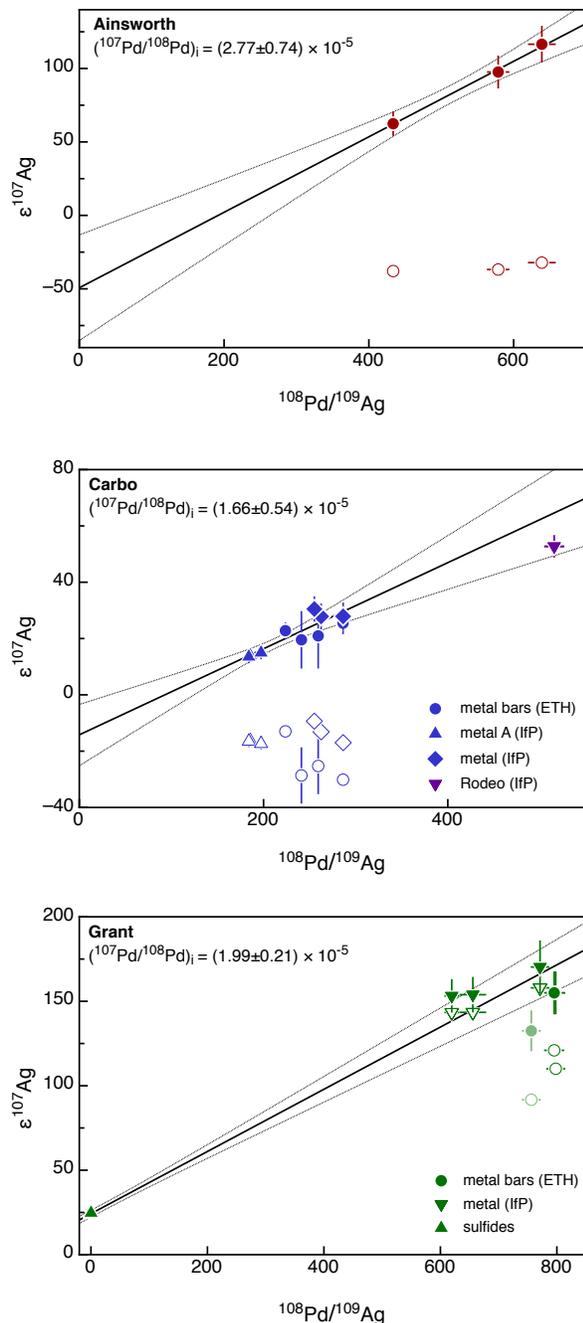


Figure 1. Pd-Ag isotope systematics. Open symbols represent measured values, GCR corrected values are shown with closed symbols. Regressions calculated using IsoPlot.

Discussion: The calculated GCR-induced shifts on $\epsilon^{107}\text{Ag}$ range from $\epsilon^{107}\text{Ag}_{\text{GCR}}$ values of *ca.* -10 for some Grant samples to values of *ca.* -150 for Ainsworth. After correction, the different metal samples from Ainsworth and Carbo, which are all characterized by negative measured $\epsilon^{107}\text{Ag}$, display positive $\epsilon^{107}\text{Ag}$ values and well-resolved ^{107}Ag excesses, as expected for their high $^{108}\text{Pd}/^{109}\text{Ag}$ (Fig 1). Thus, GCR-induced

effects on $\epsilon^{107}\text{Ag}$ not only exceed the analytical uncertainty of the Ag isotope measurements, but in case of Carbo and Ainsworth are also larger than the radiogenic ingrowth from ^{107}Pd -decay. Consequently, correction of GCR-effects is a prerequisite for obtaining accurate Pd-Ag ages for iron meteorites.

Several observations indicate that our correction method for GCR-effects on $\epsilon^{107}\text{Ag}$ is valid and provides reasonable results. First, in spite of measured negative $\epsilon^{107}\text{Ag}$, the GCR-corrected $\epsilon^{107}\text{Ag}$ values are all positive, as expected for iron meteorites with high $^{108}\text{Pd}/^{109}\text{Ag}$. Second, the GCR-corrected $\epsilon^{107}\text{Ag}$ values are positively correlated with $^{108}\text{Pd}/^{109}\text{Ag}$ for all three investigated samples (Fig. 1). Note that a significant underestimation of the GCR-induced effects on $\epsilon^{107}\text{Ag}$ would affect the samples with the highest $^{108}\text{Pd}/^{109}\text{Ag}$ the strongest, such that it would be unlikely to obtain a positive correlation of $\epsilon^{107}\text{Ag}$ with $^{108}\text{Pd}/^{109}\text{Ag}$. Third, linear regressions of the GCR-corrected Pd-Ag data yield initial $^{107}\text{Pd}/^{108}\text{Pd}$ ratios between $\sim 1.7 \times 10^{-5}$ and $\sim 2.8 \times 10^{-5}$ (Fig. 1). A significant overestimation of the GCR-effects on $\epsilon^{107}\text{Ag}$ would have resulted in (apparent) isochrons that are too steep. Note, however, that the initial $^{107}\text{Pd}/^{108}\text{Pd}$ ratios obtained after GCR-correction are all similar to previously reported values for weakly-irradiated irons [1,4], and that they also are lower than the solar system initial value. Finally, the Pd-Ag age of 4563.7 ± 1.7 Ma determined here for Grant (calculated relative to the IVA iron Muonionalusta [4] using its Pb-Pb age [9], which we corrected for U isotope variations [10]) is in very good agreement with an ^{53}Mn - ^{53}Cr of 4563.6 ± 0.7 Ma for sarcoside from Grant [11] (relative to the angrite D'Orbigny).

Conclusions: Neutron capture-induced Ag isotope variations are large and must be corrected before iron meteorites can be dated using the Pd-Ag system. The correction method presented here can be used to quantify GCR-effects on Ag isotopes, using Pt isotopes as the neutron-dose monitor. With this correction method precise Pd-Ag isochrons can be obtained, even for some of the most strongly irradiated iron meteorites.

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