

**EVALUATING N<sub>2</sub>O AS AN EXOPLANET BIOSIGNATURE: COMBINING BIOGEOCHEMICAL, PHOTOCHEMICAL AND SPECTRAL MODELS.** E. W. Schwieterman<sup>1,2,3,4</sup>, S. L. Olson<sup>1,2</sup>, C.T. Reinhard<sup>5,2</sup>, and T. W. Lyons<sup>1,2</sup>, <sup>1</sup>Department of Earth Sciences, University of California, Riverside (eschwiet@ucr.edu), <sup>2</sup>NASA Astrobiology Institute, <sup>3</sup>NASA Postdoctoral Program Fellow, <sup>4</sup>Blue Marble Space Institute of Science, Seattle, WA, <sup>5</sup>School of Earth & Atmospheric Sciences, Georgia Institute of Technology

**Introduction:** Recent studies demonstrating the possibility of ‘false positives’ for biotic O<sub>2</sub> [e.g., 1] and scenarios of low CH<sub>4</sub> and O<sub>2</sub> [2], as indicated for a billion or more years of Earth history, have increased the importance of fully characterizing alternative biosignature gases to support the search for life outside the solar system. Nitrous oxide (N<sub>2</sub>O) has been proposed as an exoplanet biosignature because its abiotic sources on Earth are small and because it has potentially observable spectral features [3]. However, the absorption bands of N<sub>2</sub>O are relatively weak, and no study has thus far examined the realistic limits of N<sub>2</sub>O detectability assuming both chemically plausible fluxes and photochemistry. Here we present preliminary results from a novel approach coupling the ocean biogeochemistry model GENIE [2] with the NAI-VPL *Atmos* photochemical model [4] to predict the mixing ratio profiles of N<sub>2</sub>O for a variety of oxygenation and flux conditions. We then model the spectral observability of these predicted N<sub>2</sub>O concentrations in simulated emission and transmission spectra using the SMART radiative transfer code [5]. We describe the limiting scenario for an ocean biosphere below.

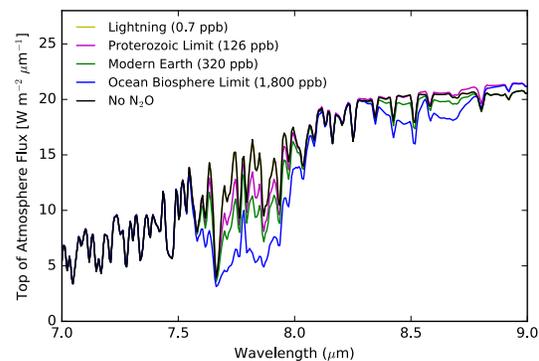
**Scenario:** N<sub>2</sub>O production in the ocean is inhibited by oxygenation of the water column, but could have been substantially greater during parts of the Proterozoic Eon (2.5-0.8 Ga) after the advent of oxygenic photosynthesis, but before the surface ocean was fully oxygenated. However, N<sub>2</sub>O is more effectively photolyzed with reduced O<sub>2</sub>, which provides a shielding effect. A recent estimate for *p*O<sub>2</sub> during the Proterozoic is 0.1% PAL, but could have been as high as 2.5% PAL without significant oxygenation of the deep ocean and reduced N<sub>2</sub>O generation [6]. An optimistic limit for the N<sub>2</sub>O flux from the ocean is obtained by assuming that nitrate is the sole electron acceptor for carbon mineralization, and that complete reduction to N<sub>2</sub> was inhibited, perhaps by trace metal limitation, i.e., all primary production results in an N<sub>2</sub>O flux. In this case, we obtain a global N<sub>2</sub>O flux of 3.2x10<sup>13</sup> mol/year.

**Results:** Inputting the above flux into the photochemical model results in a ground-level N<sub>2</sub>O concentration of 126 ppb for *p*O<sub>2</sub> = 0.1% PAL, and 1,800 ppb for *p*O<sub>2</sub> = 2.5% PAL. We find that, for a Sun-like host star, ocean-atmosphere systems will tend to produce maximum steady state concentrations of ~2 ppm N<sub>2</sub>O assuming Earth-like levels of primary productivity, or about 6X the modern N<sub>2</sub>O concentration (~320 ppb), which is mostly derived from the terrestrial biosphere.

Greater concentrations of N<sub>2</sub>O would require large contributions from a terrestrial biosphere far exceeding that of modern day Earth. For contrast, an estimate for abiotic generation of N<sub>2</sub>O from lightning is <1 ppb [7].

N<sub>2</sub>O has significant spectral bands centered at 3.7, 4.5, 7.8, and 8.6 μm. Figure 1 shows the strength of the MIR bands as a function N<sub>2</sub>O concentrations resulting from the limiting N<sub>2</sub>O flux and *p*O<sub>2</sub> values described above. While N<sub>2</sub>O has weak bands between 1.3 and 3.7 μm, we find they produce insignificant absorption for N<sub>2</sub>O concentrations plausible from an ocean biosphere. This is potentially problematic because future space-based telescope concepts such as LUVOIR will likely not have access to wavelengths longer than 2.5 μm. Additionally, photochemical destruction at high altitudes will mute the transit transmission signatures of N<sub>2</sub>O. In sum, we find that biogenic N<sub>2</sub>O would only be detectable through direct infrared (> 3.7 μm) observation at levels plausible for an ocean-based biosphere.

**References:** [1] Harman et al. (2015) *ApJ*, 812:137. [2] Olson et al. (2016) *PNAS*, 113:11447–11452. [3] Sagan, et al. (1993) *Nature* 365:715–21. [4] Arney et al. (2016) *Astrobiology* 16(11):873-899. [5] Meadows & Crisp. (1996) *JGR* 101:4595–4622. [6] Reinhard et al. (2016) *PNAS*, 113:8933–8938 [7] Schumann & Huntrieser, H. (2007) *ACPD* 7:2623–2818.



**Figure 1** – Synthetic mid-infrared (7-9 μm) emission spectrum of an Earth-like planet with N<sub>2</sub>O levels consistent with five scenarios: 1) No N<sub>2</sub>O (black), 2) production from lightning only (yellow; 0.7 ppb), 3) the Proterozoic limit assuming *p*O<sub>2</sub> = 0.1% PAL (magenta, 126 ppb), 4) the modern Earth level (green, 320 ppb), 5) and the maximum limit for an ocean biosphere, assuming (3) except *p*O<sub>2</sub> = 2.5% PAL (blue, 1,800 ppb).