GEOLOGIC MAP OF THE LACHESIS TESSERA QUADRANGLE (V-18), VENUS. D. L. Buczkowski ${ }^{1}$, L. A. Fattaruso ${ }^{2}$, E. M. McGowan ${ }^{2,3}$, and G. E. McGill ${ }^{2}$. ${ }^{1}$ Johns Hopkins Applied Physics Laboratory, Laurel, MD 20723; ${ }^{2}$ University of Massachusetts, Amherst, MA; ${ }^{3}$ Mount Holyoke College, South Hadley, MA.

Introduction: The Lachesis Tessera V-18 quadrangle ( $25^{\circ}-50^{\circ} \mathrm{N}, 300^{\circ}-330^{\circ} \mathrm{E}$ ) includes parts of Sedna and Guinevere Planitiae; regional plains [1] cover $\sim 80 \%$ of the quadrangle. The region includes 2 deformation belts and embayed fragments of 1-2 possible additional belts, 3 large central volcanoes, abundant small shield volcanoes and associated flow materials, 13 impact craters, 3 named coronae, many coronae-like features, arachnoid-like features, and dark spots [2]. The quadrangle contains a linear grouping of a prominent NW to SE oriented structural belt, coronae, and coronae-like structures [2]. Important individual structural features include radar-bright lineaments, graben, and wrinkle ridges, as well as broader ridges that may be local folds in regional plains. We present an updated version of the geologic map (Fig. 1).

Methods: Mapping was based on a $250 \mathrm{~m} / \mathrm{p}$ Magellan cycle 1 synthetic aperture radar (SAR) $1: 5 \mathrm{M}$ scale controlled mosaic, but most of the analysis utilized $75 \mathrm{~m} / \mathrm{p}$ FMAPS. Topographic information was derived from digital elevation models and from gridded elevation data; altimetry data were combined with SAR data to create synthetic stereoscopic images.

Geology: Stratigraphic units were defined primarily by their radar brightness and surface textures, although crosscutting relations, relative ages, and apparent association with topographic or structural features were also considered.

Basement materials: Tessera material (t) appears very bright on SAR images, due to being heavily deformed by at least two dominant trends of graben, ridges and penetrative lineations at high angles to each other. Tessera are sharply embayed and/or truncated by all other mapped units and are thus inferred to be the oldest material in the quadrangle. The relative ages of isolated patches of tessera cannot be determined.

In some regions, tessera material is heavily embayed by plains or volcanic units to the extent that the surface has similar areas of both units but on a scale too small to be delineated at the map scale. These units have been labeled embayed tessera (te).

Tessera-like material (tq) does not exhibit a pattern of two or more well defined deformation trends at high angles to each other; instead it is characterized by multiple deformation trends that generally are spatially distinct from each other. $\mathbf{t q}$ is radar bright, highly deformed, and extensively flooded by plains materials.

Several very small patches of radar bright, hummocky to lineated material ( $\mathbf{m b}$ ) are clearly embayed by regional plains, and thus have a relative age similar to
that of tessera. Exposures are widely scattered, and too small to determine age relations with tessera.

Plains materials: Regional plains, the most extensive materials in the quadrangle, are mapped as two units, based on radar brightness. While the number of impact craters superposed on the plains is too small to measure age differences between the emplacement of the darker (pr1) and brighter (pr2) regional plains units [3], stratigraphic markers imply that pr2 is younger than pr1. However, clear cut examples of wrinkle ridges and fractures superposed on pr1 but truncated by pr2 have not been found, indicating that the age difference is very small.

Dark plains materials (pd) are a local plains unit that are defined by significantly lower radar backscatter; they are superposed on regional plains materials, with local inliers of regional plains. Mottled plains materials ( $\mathbf{p m}$ ) consist of intermingled patches of material that is either dark or moderately bright on SAR images, resulting in a mottled pattern at $50-100 \mathrm{~km}$ scale. Its age relative to the regional plains is uncertain.

Volcanic materials: Volcano "a" materials (mva) consist of a large field of digitate flows of variable brightness (mostly moderately bright) extending to the east and northeast from the caldera. Proximal flows are smaller and younger than distal flows. Volcano "b" materials (mvb) extend generally eastward from their caldera, and consist of a field of digitate flows of variable brightness, with the brightest parts of individual flows located farthest from the caldera. Flows are superposed on regional plains materials.

Shield flows (fs, fsd) are large flow fields associated with abundant small shields. Most exposures of fs are moderately brighter than regional plains, but in one locality shield flows (fsd) are darker. Small shields similar to those within unit fs are common as isolated small groups or individuals superposed on unit pr1. Isolated flows (f) are mostly moderately bright, relatively rare, digitate flows that generally do not have a resolvable construct at their source. Pancake domes are moderately bright and delineated as unit (vd).

Corona materials: At least four coronae are present in V-18. These are widely separated, so it is not possible to determine their relative ages. The structures defining the coronae cut materials of both regional plains units. The ages of the generally sparse flows associated with the coronae relative to plains and other materials are ambiguous, although some corona materials appear to be younger than adjacent regional plains.

Zemire Corona has locally digitate flows (fcZ) that are moderately bright to bright, both within and adjacent to the corona, that include many small shields and several paterae; they are the most extensive corona flows in the quadrangle. Renenti Corona has moderately bright digitate flows (fcR) adjacent to the corona but not within it. Pasu-Ava Corona has moderately bright and homogeneous material ( $\mathbf{f c P A}$ ) with no clear flow forms that occurs within and adjacent to the corona. Jaszai Patera flows (fcJ) are moderately bright: there is a potential companion corona-like feature just NNE. There is a putative, partially imaged corona a short distance east of Pasu-Ava with no associated flows.

Deformation belts: Many individual structures occur in poorly defined belts that do not include associated mappable materials. However, there are instances where deformation belts include mappable material that can be distinguished from surrounding regional plains, although the relative ages are ambiguous. Ridge belt material (br) consist of closely spaced ridges, radarbright lineaments and, rarely, graben. Fracture belt material (bf) forms radar-bright belts defined by closely spaced graben and lineaments.

Impact crater materials: V-18 craters range from 2.4 to 40 km diameter and are superposed on either regional plains or on flows that are, in turn, superposed on regional plains. Craters are mapped as crater materials, undifferentiated (c) including materials of central peaks, walls, rims, floors and ejecta. These materials are bright to very bright. Very bright digitate flow material associated with some craters is mapped as crater flow material ( $\mathbf{c f}$ ).

Features of Interest: Important individual structural features in the V -18 quadrangle include radar-bright lineaments, graben, and wrinkle ridges, all of which are abundant and pervasive. Wrinkle ridges are sinuous, radar-bright linear features that formed approximately normal to compressive stresses in the shallow crust.

Deformation belts vary widely in trend with respect
to each other, even exhibiting significant variations within individual belts. Shishimora Dorsa is the largest and most clearly defined ridge belt, with $\sim 250 \mathrm{~km}$ within V-18. The belt trends NE and is elevated relative to adjacent regional plains. Although dominated by ridges, the belt also includes radar-bright lineaments with two distinct azimuths that define a grid pattern.

Breksta Linea is the largest fracture belt in V-18, at about 500 km long. It consists of closely spaced fractures and graben, most of which trend with about the same azimuth as the belt itself. The fractures appear to be younger than the regional plains adjacent to Breksta Linea, as they crosscut the contact between pr 1 and pr 2 in about the center of the belt. Breksta Linea is elevated relative to adjacent regional plains.

Coronae and fragments of other ridge belts extend to the southeast of Breksta Linea forming the most prominent structural grouping in the quadrangle. Coronae here are commonly flanked by pancake domes.

Three large central volcanoes have been identified, but flow units are only mapped for two of them. Topographic data strongly suggests the presence of a third large volcano in a data gore. Fractures radial to the proposed volcano are identified in the radar imagery.

Eleven named craters are located within the V-18 quadrangle. We have identified 5 other features that may also be impact craters.

We also observe $>10$ dark spots between $10-50 \mathrm{~km}$ diameter that are without topographic relief. These are not uncommon features on Venus and are often interpreted as fine debris. At least 6 of the dark spots in this quadrangle have a distinct annulus, which suggests a possible relationship to impact craters or other degraded or eroded circular surface phenomena.

References: [1] McGill G.E. (2000) V-20 quadrangle [2] McGowan E. M. \& McGill G. E. (2011) LPSC XLII, abs. 1300 [3] Campbell B.A. (1999) JGR 104, 21,951-21,955.


Figure 1. Current draft of the geologic map of the V-18 quadrangle. Areas where SAR imagery is visible are regional plains 2 unit (pr2).

