

**CO3 AND CH/CB CAIs SUGGEST  $^{10}\text{Be}$  WAS DISTRIBUTED UNIFORMLY IN THE SOLAR NEBULA.**

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**Introduction:** Studies of  $^{10}\text{Be}$ - $^{10}\text{B}$  isotope systematics in the first-formed Solar System solids (calcium-aluminum-rich inclusions, or CAIs) can provide insights into the astrophysical environment of the early Solar System. The short-lived radionuclide  $^{10}\text{Be}$  decays to  $^{10}\text{B}$  ( $t_{1/2} = 1.4$  Ma) and is produced almost exclusively by energetic particle irradiation [1]. The most likely astrophysical sites for  $^{10}\text{Be}$  production are: (1) in the molecular cloud, as galactic cosmic rays (GCRs) spalled heavier nuclei, followed by  $^{10}\text{Be}$  homogenization as the cloud collapsed to form the solar nebula [2,3]; or (2) in the nebular disk, as solar flare high-energy particles interacted with nebular gas and solids, producing  $^{10}\text{Be}$  heterogeneously in space and time [4-7]. Most normal CAIs studied thus far have been from CV3 chondrites; these have initial  $^{10}\text{Be}/^9\text{Be} \sim (6-9) \times 10^{-4}$  [5-11], although a few with higher  $^{10}\text{Be}/^9\text{Be}$  have been reported [12,13]. The few CH/CB CAIs studied so far have  $^{10}\text{Be}/^9\text{Be}$  like those in normal CV3 CAIs, but one CAI has a significantly higher  $^{10}\text{Be}/^9\text{Be} \sim 10^{-2}$  [14]. CAIs with Fractionation and Unknown Nuclear effects (FUN CAIs) and hibonites have lower  $^{10}\text{Be}/^9\text{Be} \sim (3-5) \times 10^{-4}$  [7,9,15,16]. Here we report  $^{10}\text{Be}$ - $^{10}\text{B}$  systematics in normal, pristine CAIs from CO3 and CH/CB chondrites, to better constrain how and where  $^{10}\text{Be}$  was produced.

**Methods:** Epoxy-mounted polished thick sections of CO3 chondrites Dar al Gani (DaG) 005 and DaG 027 contained six coarse-grained CAIs, and two polished mounts of the CH/CB chondrite Isheyevo contained five coarse-grained CAIs. We characterized the CAIs with the JEOL JXA-8530F electron microprobe at Arizona State University, then determined the  $^{10}\text{Be}$ - $^{10}\text{B}$  isotope systematics in these 11 CAIs using the IMS-1290 secondary ion mass spectrometer (SIMS) at UCLA. Using a 1-2 nA  $^{16}\text{O}_2^-$  primary beam (generated by a *Hyperion-II* source [17]), we pre-sputtered a  $10 \times 10 \mu\text{m}$  rastered square, then decreased the rastered area to  $5 \times 5 \mu\text{m}$  for the analysis. Secondary ion intensities were measured with multiple electron multipliers (EMs) with a mass resolving power of  $\sim 2,500$  in dynamic multi-collection mode. A NIST 614 glass was used as a standard to determine the  $^9\text{Be}/^{11}\text{B}$  relative sensitivity factor (RSF) and the  $^{10}\text{B}/^{11}\text{B}$  instrumental mass fractionation (IMF).

**Results:** *CO3 CAIs.* The six CO3 CAIs range in their size, mineralogy, texture, and shape: 300-830  $\mu\text{m}$  in size; coarse-grained texture with melilite  $\pm$  spinel, hibonite, perovskite; rounded to irregularly fragmented in shape. The  $^{10}\text{Be}$ - $^{10}\text{B}$  data for each of these CAIs define isochrons that yield a weighted average initial  $^{10}\text{Be}/^9\text{Be} = (8.4 \pm 1.6) \times 10^{-4}$  (2SE weighted, MSWD=0.3).

*CH/CB (Isheyevo) CAIs.* The five Isheyevo CAIs range in their size, mineralogy, and texture: 120-290  $\mu\text{m}$  in size; fine- to coarse-grained texture with melilite  $\pm$  hibonite, spine, grossite, and perovskite, all are rounded in shape. Taken together, the  $^{10}\text{Be}$ - $^{10}\text{B}$  data for all five CAIs yields an initial  $^{10}\text{Be}/^9\text{Be} = (10.5 \pm 2.8) \times 10^{-4}$  (MSWD=1.2).

**Discussion/Conclusions:** The six CO3 CAIs and five CH/CB CAIs measured in this study record initial  $^{10}\text{Be}/^9\text{Be}$  ratios similar to those in most normal CV3 CAIs, and all have the same  $^{10}\text{B}/^{11}\text{B}$  initial value within error. If well-behaved  $^{10}\text{Be}$ - $^{10}\text{B}$  isochrons (i.e., with MSWDs close to  $\sim 1$ ) of all previously studied normal CAIs from CV3 [5-12], CR2 [18], CH/CB [14], and CO3 chondrites are considered, they yield a weighted mean initial  $^{10}\text{Be}/^9\text{Be} = (7.0 \pm 0.3) \times 10^{-4}$  (2SE weighted). This suggests that  $^{10}\text{Be}$  was distributed homogeneously in at least those regions of the solar nebula where normal CAIs formed; in this scenario, FUN CAIs (with lower  $^{10}\text{Be}/^9\text{Be}$ ) either formed in an isotopically distinct region or formed later than normal CAIs. This further indicates that  $^{10}\text{Be}$  was likely produced in the molecular cloud, and was mostly homogenized during its collapse and formation of the solar nebula. If we consider this initial  $^{10}\text{Be}/^9\text{Be}$  value as the steady-state level generated in the molecular cloud by interactions with GCRs, the GCR ion flux  $\sim 4.6$  billion years ago in the region where our molecular cloud core was forming was about 8 times higher than present-day GCR ion flux in the solar neighborhood [3].

**Acknowledgments:** This work is supported by a NASA Earth and Space Science Fellowship (NNX16AP48H) to ETD and a NASA Emerging Worlds grant (NNX15AH41G) to MW.

**References:** [1] Davis A.M. & McKeegan K.D. (2014), *Treatise on Geochemistry* (2<sup>nd</sup> Ed.), 361. [2] Desch S.J. et al. (2004) *ApJ* 602, 528. [3] Tatischeff V. et al. (2014) *ApJ* 796, 124. [4] Jacquet E. (2019) *A&A accepted* [5] McKeegan K.D. et al. (2000) *Science* 289, 1334. [6] Sugiura N. et al. (2001) *MAPS* 36, 1397. [7] MacPherson G.J. et al. (2003) *GCA* 67, 3165. [8] Chaussidon M. et al. (2006) *GCA* 70, 224. [9] Wielandt D. et al. (2012) *ApJ* 748, 25. [10] Srinivasan G. et al. (2013) *EPSL* 374, 11. [11] Mishra R.K. & Marhas K.K. (2019) *Nature Astronomy* s41550-019-0716-0 [12] Sossi P.A. et al. (2017) *Nature Astronomy* 1, 0055. [13] Dunham E.T. et al. (2017) *Meteoritics & Planetary Science* 80:A6381 [14] Gounelle M. et al. (2013) *ApJ* 763, 33. [15] Liu M.-C. et al. (2009) *GCA* 73, 5051. [16] Fukuda K. et al. (2018) *Geochem J.* 52, 3. [17] Liu M.-C. et al. (2018) *IJMS* 1, 9 [18] Dunham E.T. et al. (2019) LPS L, Abstract #1928.