

**A VOLCANIC (QUENCHED) ANGRITE CLAST IN POLYMICT UREILITE DAG 319.** C.A. Goodrich<sup>1</sup>, T. Mikouchi<sup>2</sup> and A.H. Treiman<sup>3</sup>. <sup>1</sup>Planetary Science Institute, [goodrich@psi.edu](mailto:goodrich@psi.edu); <sup>2</sup>Univ. Tokyo; <sup>3</sup>Lunar & Planetary Institute.

**Introduction:** Typical polymict ureilites are breccias thought to have formed in a regolith layer on a ureilitic asteroid [1,2]. They consist mostly of clasts of main group ureilite materials. However, they also contain a few % non-indigenous clasts that represent impactors. These include fragments of enstatite, ordinary, carbonaceous, and Rumuruti-type chondrites [1-4], as well as angrites [3,5,6], a group of plutonic/volcanic achondrites [7]. Previously identified angrite clasts appear to be the plutonic variety. We describe a clast of a volcanic angrite in DaG 319.

**Methods:** Using whole section WDS X-ray maps for Si, Mg, Fe, Al and Ca, we identified clast B40 in section DaG 319\_M1 as having much higher Fe, Al and Ca contents, and being finer-grained, than main group ureilite materials.

**Petrography and Mineral Compositions:** Clast B40 is ~1×1.5 mm in size. It is roughly elliptical and has irregular boundaries with surrounding ureilitic matrix. The clast has a subophitic texture, consisting of subhedral to anhedral olivine grains (~5-130 μm), plagioclase laths (~10-40 μm), and intergranular pyroxene. Minor phases include kirschsteinite, Fe-Ti oxide(s), and Fe-sulfide.

Olivine compositions range from mg# 28 to 92. CaO contents of olivine are inversely correlated with mg#, and range from ~0.06 to 7.0 wt.%. Most olivine grains are normally zoned. The 4 largest (≥100 μm) have ~20-80 μm-sized cores of mg# 88-92, and are strongly zoned to mg# ~25 rims. Boundaries between cores and rims are diffuse. Smaller olivines have more ferroan cores and are less strongly zoned, but there is no significant gap in size between the 4 largest grains and smaller ones. Cr<sub>2</sub>O<sub>3</sub> contents are ~0.2-0.25 wt.% in the most magnesian cores, and much lower (≤0.06 wt.%) in rims and smaller olivines. Cores of some of the smaller olivines contain many ≤μm-sized inclusions (glass?) and/or pits.

Plagioclase is An 99. Pyroxene is rich in Ca, Al and Ti, with Wo ~51, mg# 53-59, ~7-8 wt.% Al<sub>2</sub>O<sub>3</sub> and ~1.5 wt.% TiO<sub>2</sub> (zonation of pyroxene has not yet been investigated). Kirschsteinite (~22% CaO, mg# 13-15) occurs as overgrowths on olivine.

**Discussion:** Mineralogy and mineral compositions of clast B40 are consistent with those of angrites [7]. Texture and grain sizes are within the range of those in the volcanic angrites, and are similar to the groundmass of Asuka 881371 [8,9]. The absence of a significant gap in size between the largest, most magnesian olivines and smaller ones is a significant difference from Asuka 881371, in which large olivines are interpreted as xenocrysts [8,9]. Nevertheless, the magnesian cores of the largest olivines in clast B40 may be remnants of partially resorbed xenocrysts, as interpreted for the similar-sized olivines in LEW 87051 [10]. This clast adds to the array of cooling histories of volcanic angrites.

**References:** [1] Goodrich C.A. et al. 2004. *Chemie der Erde* 64, 283-327. [2] Downes H. et al. 2008. *GCA* 72, 4825-4844. [3] Ikeda Y. et al. 2000. *Ant. Met. Res.* 13, 177-221. [4] Goodrich C.A., Gross J. 2015. *LPSC* #1214. [5] Prinz M. et al. 1986. *LPS* 17, 681. [6] Jaques A.L., Fitzgerald M.J. 1982. *GCA* 46, 893-900. [7] Keil K. 2012. *Chemie der Erde* 72, 191-218. [8] McKay G. et al. 1995. *MAPS* 30, 543. [9] Mikouchi T. et al. 1995. *MAPS* 30, 549. [10] Mikouchi T. et al. 1996. *Proc. NIPR Symp. Ant. Met.* 9, 174-188.