## FORMATION PERIOD FOR A FLUFFY TYPE A CAI FROM VIGARANO.

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Introduction: Fluffy Type A CAIs containing reversely zoned melilite crystals are suggested to be direct condensates from a solar nebular gas [1]. The O isotope measurements with a line-profile technique using SIMS for single reversely zoned melilite crystals in a fluffy Type A CAI, V2-01 from Vigarano CV3 have revealed the O isotopic compositions change from <sup>16</sup>Opoor to <sup>16</sup>O-rich along with their crystal growths, indicating that the O isotopic compositions of surrounding nebular gas were changed from <sup>16</sup>O-poor to <sup>16</sup>O-rich during their condensations [2]. Such trends have also been found in other Type A CAIs from Allende and Efremovka [3, 4]. In this study, we have measured O isotopic compositions of spinel enclosed by melilite crystals, and spinel and diopside from the Wark-Lovering rim (W-L rim) in the V2-01 CAI, to understand the O isotope change during the total formation of V2-01. We have also investigated <sup>26</sup>Al-Mg systematics of individual reversely zoned melilite crystals, spinel, and diopside in V2-01. The O and Al-Mg isotope measurements were conducted using SIMS of Hokkaido University (Cameca ims-1280HR). Time duration for a formation of V2-01 and O isotope change during the formation have been determined.

**Results and discussion:** The spinel enclosed by melilite crystals showed <sup>16</sup>O-rich compositions ( $\Delta^{17}$ O ~ -24‰). The spinel and diopside from W-L rim also had homogeneously <sup>16</sup>O-rich compositions ( $\Delta^{17}$ O ~ -23‰). On the other hands, the O isotopic compositions of surrounding nebular gas during the formation of reversely zoned melilite crystals have been suggested to be changed from <sup>16</sup>O-poor ( $\Delta^{17}$ O ~ -5‰) to <sup>16</sup>O-rich ( $\Delta^{17}$ O ~ -25‰) [2]. The petrography suggests that formation sequences of these minerals were spinel enclosed by melilite, reversely zoned melilite crystals, and spinel and diopside in the W-L rim. It indicated that V2-01 was formed in a variable O isotope reservoir changing from <sup>16</sup>O-rich to <sup>16</sup>O-poor and back to <sup>16</sup>O-rich.

The <sup>26</sup>Al-Mg systematics is consistent to the formation sequences. The spinel enclosed by melilite showed an initial value of  $({}^{26}\text{Al}/{}^{27}\text{Al})_0 = (5.7 \pm 0.2) \times 10^{-5}$ . The six reversely zoned melilite crystals measured showed indistinguishable initial values of  $({}^{26}\text{Al}/{}^{27}\text{Al})_0 \sim 4.7 \times 10^{-5}$ . The spinel and diopside from W-L rim showed an initial value of  $({}^{26}\text{Al}/{}^{27}\text{Al})_0 = (4.6 \pm 0.4) \times 10^{-5}$ , which is comparable to the initial values of melilite crystals. This is a first determination of a formation age of the W-L rim. The time duration for the V2-01 CAI formation could be estimated as 0.22  $\pm$  0.11 Myr from the difference of initial values. Therefore, our study revealed that the V2-01 was formed within the time duration of 0.22  $\pm$  0.11 Myr in the variable O isotope reservoir changing from <sup>16</sup>O-rich to <sup>16</sup>O-poor and back to <sup>16</sup>O-rich.

**References:** [1] MacPherson G. J. and Grossman L. 1984. *Geochimica et Cosmochimica Acta* 48: 29–46. [2] Katayama J. et al. 2012. *Meteoritics and Planetary Science* 47: 2094–2106. [3] Kawasaki N. et al. 2012. *Meteoritics and Planetary Science* 47: 2084–2093. [4] Park C. et al. 2012. *Meteoritics and Planetary Science* 47: 2070–2083.