Temperature and Pressure Effects on Diffusional Mg Isotopes in MgSiO₃ Liquid. Haiyang Luo¹, Bijaya B. Karki², Dipta B. Ghosh², Xiaobin Cao¹, Huiming Bao¹, ¹Department of Geology & Geophysics, Louisiana State University, Baton Rouge, LA 70803, USA, ²Department of Computer Science, Department of Geology and Geophysics, Center for Computation and Technology, Louisiana State University, Baton Rouge, LA 70803, USA

Introduction: Diffusional isotope fractionation (DIF) is an important parameter when disequilibrium systems are of interest, e. g. during crystal growth from a silicate melt. DIF was found to be large even at silicate melt temperatures[1,2,3,4]. Unlike simple diffusion of a molecular species in a dilute gas, it has been difficult to predict diffusional isotope fractionation factors for melt systems. Usually, β factor, a dimensionless empirical parameter, is used to describe the mass dependent diffusional isotope fractionation. The β factor is often assumed to be dependent on variables such as temperature, pressure, and chemical composition. However, the degrees of these variables' effect on the β factor are not explored. Here, we used first-principles molecular dynamic simulations to explore diffusional isotope effect of Mg in MgSiO₃ melt under a series of pressures (0 – 68 GPa) and temperatures (3000, 4000, 5000 K).

Results and Discussion: Our results show that, at a given temperature the β factor for diffusional Mg isotope effects decreases when pressure increases (Fig. 1). At 5000 K, β decreases from 0.270±0.004 to 0.164±0.029 with pressure increasing from 6 to 68 GPa (Fig. 2). At 4000 K, β decreases from 0.255 \pm 0.007 to 0.182±0.040 with pressure increasing from 1 to 60 GPa. At 3000 K, β decreases from 0.288±0.014 0.239±0.024 with pressure increasing from 0 to 10 GPa. We can attribute this decrease to the chemical environmental change of Mg in MgSiO₃ liquid at atomic level. At a constant volume, no clear trend is observed between the β factor and temperature, indicating that the temperature effect on the chemical environmental change of Mg is limited under the temperature range we explored.

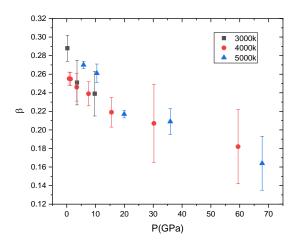


Fig. 1 β versus pressure at 3000, 4000, 5000 K.

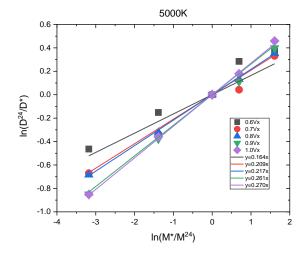


Fig. 2 $\ln(D^{24}/D^*)$ versus $\ln(M^*/M^{24})$.

References: [1] Richter FM et al. (1999) *GCA*, *63*, 2853–2861. [2] Richter FM et al. (2003) *GCA*, *67*, 3905–3923. [3] Richter FM et al. (2008) *GCA*, *72*, 206–220. [4] Watkins JM et al. (2009) *GCA*, *73*, 7341–7359.