

**PLUTO AND CHARON: SURFACE COLORS AND COMPOSITIONS--A HYPOTHESIS** D. P Cruikshank<sup>1</sup>, S. A. Stern<sup>2</sup>, W. M. Grundy<sup>3</sup>, J. M. Moore<sup>1</sup>, L. A. Young<sup>2</sup>, C. B. Olkin<sup>2</sup>, H. A. Weaver<sup>4</sup>, K. Ennico<sup>1</sup>, J. C. Cook<sup>2</sup>, and the New Horizons Composition Theme Team <sup>1</sup>NASA Ames Research Center (Dale.P.Cruikshank@nasa.gov), <sup>2</sup>Southwest Research Institute, <sup>3</sup>Lowell Observatory, <sup>4</sup>Appl. Phys. Lab Johns Hopkins Univ.

**Introduction:** The surface of Pluto displays an array of colors ranging from yellow to red to brown, while the surface of Charon is largely gray with a north polar zone of red color similar to regions on Pluto[1,2]. Pluto's surface shows layers of intensely colored material in tilted and transported blocks, and fractured geographical units. This arrangement suggests episodes of formation or deposition of that material interspersed with episodes of emplacement of ices having little or no color. The ices identified on the surfaces of these two bodies (N<sub>2</sub>, CH<sub>4</sub>, CO, C<sub>2</sub>H<sub>6</sub>, H<sub>2</sub>O on Pluto, and H<sub>2</sub>O and NH<sub>3</sub> on Charon)[3,4] are colorless, as are nearly all ices in a powdery state. The colors on Pluto probably arise from the *in situ* formation of a macromolecular carbonaceous material generated by energetic processing of the ices on the surface[5]. Laboratory experiments producing refractory tholins particularly relevant to Pluto explored the chemistry of both UV and low-energy electron bombardment of a mix of Pluto ices (N<sub>2</sub>:CH<sub>4</sub>:CO = 100:1:1)[6,7]. We can term this Pluto ice tholin PIT.

Water ice in the crystalline state characterizes Charon's surface, and while most of Charon's surface is neutral in color, with geometric albedo ~0.38 [2,3], the polar zone and a light cover of fainter but similar reddish color over some surface regions suggest a common origin with the colored material on Pluto. NH<sub>3</sub> or NH<sub>3</sub>•nH<sub>2</sub>O was identified from disk-integrated Earth-based spectra[e.g.,8,9], and a few concentrated NH<sub>3</sub> exposures have been found in the New Horizons spectral images [10].

**Open Questions:** In terms of the composition and evolution of the surfaces of Pluto and Charon, a number of puzzles present themselves. Why does H<sub>2</sub>O ice dominate the surface of Charon? What is the origin of the NH<sub>3</sub> or NH<sub>3</sub> hydrate found in low abundance over most of the surface, with a few isolated outcrops of higher concentration? How can NH<sub>3</sub> persist on Charon when it is readily destroyed by UV radiation? Why is the overall albedo of Charon rather low and the color mostly neutral, except for the north polar region, which displays a color similar to many regions on Pluto? Why is NH<sub>3</sub> or its hydrate not identified on Pluto, while a number of other molecules are?

**Toward an Understanding:** Turning first to recent laboratory studies of the refractory residues left from the UV or electron irradiation of the Pluto ice mix (PIT) noted above, chemical analysis shows the pres-

ence of a number of small molecules, including NH<sub>3</sub>, plus complex macromolecular carbonaceous material. Mass spectra of organic molecules desorbed from the residue show a high degree of aromatization, while UV fluorescence imaging shows an overall high concentration of organics. The yellow-to-brown colors of the residue are consistent with the presence of conjugated carbon-carbon chains. Material of comparable complexity characterizes the insoluble organic matter (IOM) in carbonaceous meteorites[11]. IOM is found to contain small molecules and radicals trapped in the carbonaceous complex for astronomical time scales. The PIT also contains small molecules, notably NH<sub>3</sub>, and probably radicals that are stable on long time-scales.

The low albedo and neutrally colored regions of Charon may be caused by the presence of tholin, possibly PIT, that has largely been radiation processed to the point where most of the hydrogen is lost and the color has consequently vanished. This over-processing, or carbonization of tholins is known from early work[e.g., 12]. Such over-processing would be expected for an old surface that is not refreshed with PIT, which might either be excavated from depth, where it was made in a previous epoch of this body's history, or made currently or recently from an inventory of N<sub>2</sub> and CH<sub>4</sub> ices. In this view, the surface underlying the gray material bears orange PIT that has been protected from overprocessing, or possibly a mix of CH<sub>4</sub> and N<sub>2</sub> ices in which PIT has been made relatively recently through the same or similar processes as are currently active on Pluto itself. The visible surface ice layer(s) of Charon are dominantly composed of the very refractory crystalline H<sub>2</sub>O ice, as noted above, and its low albedo is a consequence of intermixing or light blanketing by carbon-dominated, over-processed tholin.

Concerning the current presence of NH<sub>3</sub> or NH<sub>3</sub>•H<sub>2</sub>O exposed over most of Charon, NH<sub>3</sub> synthesized in the tholin-making process and trapped for a long time can slowly diffuse out of the complex tholin matrix, diffuse through the overlayer of H<sub>2</sub>O ice, and appear in small concentration on the optical surface, as detected spectroscopically. The destruction by UV and replenishment by upward diffusion would, in principle, reach some balance, and that balance could be different in different epochs of the evolution of the Pluto-Charon system. It is also noteworthy that NH<sub>3</sub> diffusing

through H<sub>2</sub>O results in the hydration of the ammonia, thus producing NH<sub>3</sub>•nH<sub>2</sub>O[13,14]. The hydration is facilitated if the ice structure is disrupted, as can occur by sputtering or micrometeoroid impacts. Other small molecules may also diffuse out of the tholin, but either fail to emerge on the surface with the H<sub>2</sub>O ice, or go undetected because of low concentration or spectroscopic insensitivity. Other such molecules could include H<sub>2</sub>CO, HCOOH, and N-bearing species. There has been no detection of NH<sub>3</sub> or its hydrate(s) on Pluto to date, but they may be present and undetectable because of low concentration and the interference of other, stronger absorption bands of the known constituents.

In order for tholin to be made on Charon's surface, the necessary ingredients must be present, at least for some period of time. A nitrogen-bearing molecule is clearly needed to produce the NH<sub>3</sub> now seen. PIT is N-rich (N/C~0.9 for the electron-generated tholin[6]), so the very volatile N<sub>2</sub> is needed, in addition to CH<sub>4</sub>, which is significantly less volatile. The presence of colored PIT or similar material on Charon is indicated by the coloration of the north polar region. Thus, while a primordial inventory of NH<sub>3</sub> could be native to Charon, it appears more likely that N<sub>2</sub> was an initial component of the body, as it apparently was for Pluto. If, then, Pluto and Charon shared a common inventory of primordial ices, there could have been an epoch in Charon's history when a PIT-like material would form there before all the N<sub>2</sub> was lost. Remaining CH<sub>4</sub> ice would be radiation processed to the point of the loss (or near loss) of the CH<sub>4</sub> spectral signature and its conversion to elemental carbon, which now appears as a neutral, albedo-lowering component of the H<sub>2</sub>O ice surface.

This view of the chemical evolution of Charon's surface appears to require that some amount of N<sub>2</sub> and CH<sub>4</sub> survived the event of its formation so that the raw materials for tholin formation were available for the synthesis of the colored and neutral components. Pluto's volatile inventory somehow survived the moon-forming event, as well as tidal heating that must have occurred after Charon was formed and as the two bodies became locked in rotation-revolution. Similarly, some fraction of Charon's volatiles must have survived the tidal heating.

**Summary and Conclusions:** A hypothesis is presented to link the color and chemical characteristics of Pluto and Charon to the histories of their volatile inventories and to recent laboratory experiments. Pluto's surface and atmosphere are chemically active in that photolysis and radiolysis of N<sub>2</sub>, CH<sub>4</sub>, and CO found in both the gaseous and solid states are in progress, producing refractory materials that color the landscape to varying degrees. The exact composition and rate of

formation of the coloring agent(s) are most likely dependent on the ages of surface units and the atmospheric density and composition over time.

NH<sub>3</sub> found on Charon may be explained as a product of the radiation processing of the same simple molecules as now found on Pluto and sequestered in the resulting complex organic material that is referenced here as Pluto ice tholin (PIT). While the modern surface of Charon is denuded of N<sub>2</sub> and (possibly) CH<sub>4</sub> ices, the PIT generated in an earlier epoch underlies the H<sub>2</sub>O ice surface. NH<sub>3</sub> diffuses out of the PIT and through the H<sub>2</sub>O to emerge on the surface, where it soon is lost by dissociation. Some or all of the NH<sub>3</sub> is hydrated by the process of diffusion through H<sub>2</sub>O ice.

The colored region at Charon's north pole results from the excavation of PIT formed in an earlier epoch by the removal of the H<sub>2</sub>O-ice-dominated overlayer. The region around the crater informally called Organa[10] is enriched in NH<sub>3</sub> by some aspect of the impact process that produced it, perhaps by reducing the overburden of H<sub>2</sub>O ice, thus exposing a layer in which NH<sub>3</sub> that diffused out of tholin had accumulated.

The hypothesis presented here challenges scenarios of the formation of the Pluto-Charon binary and concepts of volatile retention by both components in the face of the energetics of formation and tidal locking. It further invites a closer look at the formation and evolution of the coloring components of the surfaces of both bodies.

Counter arguments to the hypothesis present here include (but are not limited to) the means by which the present-day H<sub>2</sub>O ice cover on Charon came to overlie an ancient tholin. Ejecta from impacts penetrating to the H<sub>2</sub>O ice mantle may have been effective in this sense. It is also noted that fractured blocks seen in press release images of Charon's surface do not expose any clear layer(s) of colored material.

#### References:

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