

**IMPACTS AND CRATERING ON PLUTO: IMPLICATIONS FOR THE SOURCE OF PLUTO'S NITROGEN (N<sub>2</sub>).** Kelsi N. Singer and S. Alan Stern. Southwest Research Institute, 1050 Walnut Street, Boulder, CO 80302, United States (ksinger@boulder.swri.edu, astern@swri.edu).

**Introduction:** What will Pluto and Charon's craters tell us? In preparation for the closest approach of NASA's New Horizons spacecraft on July 14<sup>th</sup>, 2015, we consider models that will relate the observed craters to Kuiper belt object (KBO) impactor populations, surface properties, subsurface structure, cratering physics/scaling, and surface-atmosphere interactions. Craters on Pluto and Charon occur under unique conditions: low impact speeds (average  $\sim 2 \text{ km s}^{-1}$  [e.g., 1,2]) an icy surface with a combination of N<sub>2</sub>, CH<sub>4</sub>, and/or H<sub>2</sub>O ices, and gravity and escapes speeds intermediate to those of the smaller and larger icy satellites. Thus craters on Pluto and Charon will provide a key comparative study to those on other worlds. Additionally, constraints derived from Pluto system craters on the size distribution of small KBOs will improve our estimates of cratering rates and refine age estimates on other solar system bodies [e.g., 3-5].

**Predicted Impactor Populations** Using observations of current KBO densities [6], typical KBO impact velocities onto Pluto of  $\sim 1\text{-}2 \text{ km s}^{-1}$  [e.g., 2], and Pluto's cross section, Stern et al [7] estimate  $\sim 14,000$  comets 1 km in diameter or larger would impact Pluto over 4 billion years. Bierhaus and Dones [2] considered different KBO subpopulations and estimated 1840 impactors  $> 1 \text{ km}$  in diameter (nominal case, 5600 for maximum case). The largest impactor predicted to hit Pluto in the last  $\sim 4$  billion years is  $\sim 60 \text{ km}$  in diameter, and on Charon  $\sim 25 \text{ km}$  [2].

We note however that that there is uncertainty in the slope of the size distribution for KBOs smaller than  $100 \text{ km}$  in diameter, which New Horizons crater data from Pluto and Charon will help constrain [see discussion in 2,5,8,9]. Additionally a correction factor of  $\sim 2\text{-}10$  may be needed [5,8] to account for collisional and dynamical erosion of the Kuiper Belt leading to a smaller number of current KBOs compared with early in Solar System history [e.g., 10-12].

**What craters will New Horizons be able to observe?** Table 1 lists the highest resolutions achieved by the New Horizons LOng Range Reconnaissance Imager (LORRI) for Pluto and Charon. The absolute best resolutions will occur over two narrow strips of terrains, thus we show in Table 1 a second number that is given for the highest resolution covering the majority of the lit hemisphere. Using a practical four-pixel minimum for crater identification, the size of the smallest craters that could be observed and other quantities are estimated in Table 1.

**Cometary Impactors, Cratering, and Pluto's N<sub>2</sub>**  
**Atmospheric Loss:** In addition to Pluto's surface ice composition being dominated by molecular nitrogen [16-17], Pluto's atmosphere consists of a  $>90\%$  mole fraction N<sub>2</sub> [18] with surface pressures estimated on the order of  $\sim 10 \text{ }\mu\text{bars}$  [19]. Models predict the N<sub>2</sub> escape rate is  $10^{27}\text{-}10^{28} \text{ molecules s}^{-1}$ , varying by an order of magnitude over the course of Pluto's year [e.g., 20]. For reference, this leads to  $\sim 1\text{-}2 \times 10^{12\text{-}13} \text{ g}$  of N<sub>2</sub> lost per year; for reference, the estimated current global atmospheric mass of Pluto, based on pure N<sub>2</sub> of  $10 \text{ }\mu\text{bars}$  and  $35 \text{ K}$  is  $\sim 5 \times 10^{17} \text{ g}$  [7].

Using the range of escape rates stated above, a simple linear extrapolation of the escape rate yields a total mass of N<sub>2</sub> lost over four billion years of  $6 \times 10^{21\text{-}22} \text{ g}$ . This is equivalent to a condensed N<sub>2</sub> surface layer on Pluto of depth  $0.4\text{-}4 \text{ km}$  [7].

**N<sub>2</sub> Delivery by Comets:** Stern et al. [7] estimate an N<sub>2</sub> mass of  $5 \times 10^{10\text{-}11} \text{ g}$  in a  $1 \text{ km}$  comet with a 50:50

Resolutions (best overall and hemispheric coverage)	4-Pixel Final Crater Diameter	Excavation Depth <sup>a</sup>	Transient Crater Diameter <sup>b</sup>	Impactor Diameter <sup>c</sup>	Number of Impacts this Size or Greater in 4 Gyr <sup>d</sup>	Min Size of Primary needed to create Secondary <sup>e</sup>
<b>Pluto</b>						
Best: $\sim 70 \text{ m px}^{-1}$	$\sim 300 \text{ m}$	$30 \text{ m}$	$250 \text{ m}$	$20 \text{ m}$	$2.4 \times 10^6$	$6 \text{ km}$
Hemi: $\sim 300 \text{ m px}^{-1}$	$\sim 1.2 \text{ km}$	$120 \text{ m}$	$1 \text{ km}$	$110 \text{ m}$	$1.1 \times 10^5$	$24 \text{ km}$
<b>Charon</b>						
Best: $\sim 150 \text{ m px}^{-1}$	$\sim 600 \text{ m}$	$60 \text{ m}$	$500 \text{ m}$	$35 \text{ m}$	$1.8 \times 10^5$	$12 \text{ km}$
Hemi $\sim 600 \text{ m px}^{-1}$	$\sim 2.4 \text{ km}$	$240 \text{ m}$	$2 \text{ km}$	$200 \text{ m}$	$7.5 \times 10^3$	$48 \text{ km}$

**Table 1:** <sup>a</sup>Excavation depth taken as  $\sim 1/10^{\text{th}}$  the transient crater diameter [13]. <sup>b</sup>For simple craters ( $D_{\text{final}} < 4 \text{ km}$  in diameter [4]), we used the apparent diameter as an approximation for the transient, taken as  $0.83 \times D_{\text{final}}$  based on lunar data [14]. <sup>c</sup>Scaling from Holsapple [15] with updated parameters from <http://keith.a.washington.edu/craterdata/scaling/theory.pdf>. Scaling for ice-on-ice impacts,  $2 \text{ km s}^{-1}$  impact speeds,  $45^\circ$  impactor,  $d/D = 0.2$  for simplicity, Pluto's gravity is  $0.66 \text{ m s}^{-2}$ , Charon's is  $0.23 \text{ m s}^{-2}$ . <sup>d</sup>Bierhaus and Dones [2] nominal rates from their Table 8. <sup>e</sup>Size of primary crater necessary to produce a secondary crater at the 4-pixel final crater diameter (based on the largest secondaries being  $\sim 5\%$  the size of the primary [13]).

ratio of H<sub>2</sub>O to refractories and typical cometary nitrogen abundances [21,22]. Using the ~14,000 comets > 1 km in diameter estimate cited above and calculating based purely on 1 km-diameter comets, this number of impacts would deliver  $\sim 7 \times 10^{14-15}$  g of N<sub>2</sub> over 4 Ga, far, far less than the anticipated N<sub>2</sub> inventory lost over time to escape.

Thus, even given uncertainties in the impactor flux, it does not appear that comets could deliver enough N<sub>2</sub> to supply Pluto's atmosphere. This is a key conclusion of our work. We note that our 14,000 impactor estimate is comparable to, but on the higher end of other predictions in the literature [1-3,7,8]. Except in the extremely unlikely scenario that a recent, relatively large impact is supplying the current atmosphere, these calculations imply endogenic sources must dominate replenishment of Pluto's N<sub>2</sub>.

**N<sub>2</sub> Excavation by Impacts:** Impacts could punch through any lag deposits of the more refractory but considerably less abundant CH<sub>4</sub> (or water-ice/other involatile materials) that one expects to build up during Pluto's history as N<sub>2</sub> escapes to space; this would allow excavated N<sub>2</sub> to resupply Pluto's atmosphere, defeating the involatile lag deposit's tendency to choke off N<sub>2</sub>. A simple estimate of N<sub>2</sub> excavated based on 1/2 the transient crater's volume is presented in Table 2 (examples per impactor size). Dividing by the mean time between impacts of a given size yields a rough approximation for the mass of N<sub>2</sub> available to resupply the atmosphere per year, per impact size. Integrating these values over impactors ranging from 0-60 km in diameter, we arrive at a total mass of  $1 \times 10^{15}$  g yr<sup>-1</sup>. This order of magnitude estimate is larger than the predicted escape rate ( $1.4 \times 10^{12-13}$  g yr<sup>-1</sup>), and thus it is possible that the cratering process can resupply Pluto's

atmosphere without need for deep-seated cryovolcanism or another tectonic or geodynamic means to resupply the atmosphere against involatile lag deposit build up from the escape process.

**Conclusions:** Comets deliver six-to-eight orders of magnitude less N<sub>2</sub> than is necessary to support the current atmospheric escape rates over Pluto's history. However, N<sub>2</sub> excavated during these impacts, might be able to resupply Pluto's atmosphere if the surface N<sub>2</sub> layer is relatively thick.

**References:** [1] Zahnle K. et al. (2003) *Icarus* 163, 263-289. [2] Bierhaus E.B. and Dones L. (2015) *Icarus* 246, 165-182. [3] de Elía G.C. et al. (2010) *A&A* 521, 23 [4] Moore J.M. et al. (2015) *Icarus* 246, 65-81. [5] Greenstreet S. et al. *Icarus* submitted. [6] Schlichting H.E. et al. (2012) *Astrophys J.* 761, 150-159. [7] Stern S.A. et al. (2015a) *Icarus* 246, 298-302. [8] Durda D.D. and Stern S.A. (2000) *Icarus* 145, 220-229. [9] Stern S.A. et al. (2015b) *Icarus* 250, 287-293. [10] Stern S.A. and Colwell J.E. (1997) *Astrophys J.* 490, 879-882. [11] Weissman P.R. and Levison H.F. (1997) In: Stern, S. A., Tholen, D. J., (Eds.), Pluto and Charon, UAP, Tucson, 559-604. [12] Farinella P. et al. (2000) In: Mannings, V., Boss, A. P., Russell, S. S., (Eds.), Protostars and Planets IV. UAP, Tucson, 1255-1282. [13] Melosh H.J. (1989) *Impact Cratering*. OUP, New York. [14] Pike R. J. (1977) *LPSC Proc.* 8, pp. 3427-3436. [15] Holsapple K.A. (1993) *Annu. Rev. Earth Planet Sci.* 21, 333-373. [16] Owen T.C. et al. (1993) *Science* 261, 745-748. [17] Cruikshank D.P. et al. (1997) In: Stern, S. A., Tholen, D. J., (Eds.), Pluto and Charon. UAP, Tucson, 221-267. [18] Yelle R.V. and Elliot J.L. (1997) In: Stern, S. A., Tholen, D. J., (Eds.), Pluto and Charon. UAP, Tucson, 347-390. [19] Lellouch E. et al. (2009) *A&A* 495, L17-L21. [20] Zhu X. et al. (2014) *Icarus* 228, 301-314. [21] Crovisier J. (1994) In: Milani, A., di Martino, M., Cellino, A., (Eds.), Asteroids, Comets, Meteors, Vol. 160, 313-326. [22] Crovisier J. (2006) *Faraday Discussions* 133, 375-385. [23] McKinnon W.B. and Schenk P.M. (1995) *GRL* 22, 1829-1832.

Impactor Diameter [km]	Transient Crater Diameter <sup>a</sup> [km]	~Final Crater Diameter <sup>b</sup> [km]	Excavated N <sub>2</sub> Mass <sup>c</sup> [g]	$\tau_{\text{impact}}$ (mean time between impacts) <sup>d</sup> [yr]	N <sub>2</sub> Avail. / yr (Column 4/ $\tau_{\text{impact}}$ ) [g N <sub>2</sub> / yr]	$\tau_{\text{N}_2 \text{ Loss}}$ (Column 4/ N <sub>2</sub> escape rate) [yr]
0.01	0.15	0.18	$1.4 \times 10^{11}$	461	$3.0 \times 10^8$	0.1
0.05	0.54	0.65	$6.0 \times 10^{12}$	8,860	$6.8 \times 10^8$	4.3
0.5	3.3	3.9	$1.4 \times 10^{15}$	608,600	$2.2 \times 10^9$	967
1	5.6	6.8	$6.9 \times 10^{15}$	$2.2 \times 10^6$	$3.2 \times 10^9$	4,938
5	20	32	$3.1 \times 10^{17}$	$4.2 \times 10^7$	$7.3 \times 10^9$	217,866
10	34	59	$1.6 \times 10^{18}$	$1.5 \times 10^8$	$1.0 \times 10^{10}$	$1.1 \times 10^6$
20	59	108	$8.0 \times 10^{18}$	$5.3 \times 10^8$	$1.5 \times 10^{10}$	$5.7 \times 10^6$
40	102	197	$4.1 \times 10^{19}$	$1.9 \times 10^9$	$2.1 \times 10^{10}$	$2.9 \times 10^7$
60	140	280	$1.1 \times 10^{20}$	$4.0 \times 10^9$	$2.6 \times 10^{10}$	$7.5 \times 10^7$

**Table 2:** <sup>a</sup>Scaling as in Table 1, footnote d, for Pluto. <sup>b</sup>Simple crater conversion from transient to final as described in Table 1, complex craters (D > 4 k) use McKinnon and Schenk [23; their eqn 1]. Although the latter was derived for the higher gravity bodies Ganymede and Callisto, there is currently no equivalent for lower gravity icy bodies. <sup>c</sup>Half of the volume of the transient crater, using a thick, 99% N<sub>2</sub> surface layer. <sup>d</sup>Bierhaus and Dones [2], nominal rates from their Table 8.