

**$^{26}\text{Al}$ - $^{26}\text{Mg}$  systematics of chondrules: Progresses and issues from the last 5 years.** K. Nagashima<sup>1</sup>, N. T. Kita<sup>2</sup>, and T.-H. Luu<sup>3</sup>, <sup>1</sup>Hawai'i Institute of Geophysics and Planetology, University of Hawai'i at Mānoa, Honolulu, HI 96822, USA (kazu@higp.hawaii.edu), <sup>2</sup>Department of Geoscience, University of Wisconsin-Madison, Madison, Wisconsin 53706, USA, <sup>3</sup>School of Earth Sciences, University of Bristol, BS8 1RJ, UK.

**Introduction:** The short-lived radionuclide decay system  $^{26}\text{Al}$ - $^{26}\text{Mg}$  (half-life  $\sim 0.7$  Myr) has been considered to be a high precision chronometer to date processes in the protoplanetary disk. Since the first report of excesses of  $^{26}\text{Mg}$  due to *in situ* decay of  $^{26}\text{Al}$  in a chondrule [1],  $^{26}\text{Al}$ - $^{26}\text{Mg}$  chronometer have been applied to chondrules from several chondrite groups. If their initial  $^{26}\text{Al}/^{27}\text{Al}$  ratios ( $(^{26}\text{Al}/^{27}\text{Al})_0$ ) represent the timing of their formation,  $^{26}\text{Al}$  ages of chondrules provide important cosmochemical constraints such as lifetime of the protoplanetary disk, chondrule-forming processes, and chondrite accretion processes. The  $^{26}\text{Al}$ - $^{26}\text{Mg}$  systematics of chondrules and implications to the evolution of protoplanetary disk have been summarized based on the data obtained before 2012 [2]. This abstract serves as a summary for the recent progresses and issues on  $^{26}\text{Al}$ - $^{26}\text{Mg}$  systematics of chondrules from the last 5 years.

**$^{26}\text{Al}$  abundances in chondrules inferred from internal isochrons:** The  $(^{26}\text{Al}/^{27}\text{Al})_0$  in chondrules from three ordinary and carbonaceous chondrite groups (LL3.0, CO3.0, and Acfer 094 ungrouped C3) are  $\sim 6\text{--}7 \times 10^{-6}$  [2]. Since then, reliable  $(^{26}\text{Al}/^{27}\text{Al})_0$  have been obtained for chondrules from CR2-3 [3-5], CH3 [6], and CV3.1 [7]. The  $(^{26}\text{Al}/^{27}\text{Al})_0$  from CR2-3 and CH3 chondrules are distinctly different from those in LL3.0 and CO3.0. While a few chondrules have  $(^{26}\text{Al}/^{27}\text{Al})_0$  of  $\sim 6 \times 10^{-6}$ , most of them are systematically lower than  $3 \times 10^{-6}$ . The  $(^{26}\text{Al}/^{27}\text{Al})_0$  of chondrules in Kaba (CV3.1) have  $(4.8 \pm 1.1) \times 10^{-6}$ , similar to/slightly lower than those of LL3.0 and CO3.0. The  $(^{26}\text{Al}/^{27}\text{Al})_0$  of the chondrules from the chondrite groups decrease in the order of LL  $\sim$  CO  $\sim$  Acfer 094  $\geq$  CV  $\geq$  CR  $\sim$  CH. In addition, each chondrite group may have multiple populations of chondrules [4].

**$^{26}\text{Al}$  abundances for chondrule precursors inferred from bulk chondrules:** The model  $(^{26}\text{Al}/^{27}\text{Al})$  ratios of chondrule "precursors" have been estimated from bulk chondrules  $^{26}\text{Al}$ - $^{26}\text{Mg}$  isotope analyses that range from  $\sim 5 \times 10^{-5}$  to  $\sim 1 \times 10^{-5}$  [8,9]. Luu et al. [9] suggested the minimum value ( $1.2 \times 10^{-5}$ ) corresponds to the time that formation of the precursors stopped or were separated from a nebular reservoir. The difference of  $(^{26}\text{Al}/^{27}\text{Al})_0$  between the bulk and *in situ* data may correspond to a time difference between the precursor formation and the last melting of chondrules.

**Homogeneous/heterogeneous distribution of  $^{26}\text{Al}$  in the disk:** To convert  $(^{26}\text{Al}/^{27}\text{Al})_0$  values to relative  $^{26}\text{Al}$  ages, we have to assume that  $^{26}\text{Al}/^{27}\text{Al}$  was homogeneously distributed throughout the disk. This assumption has been challenged and is highly controversial. Recent studies on multiple chronometers ( $^{26}\text{Al}$ - $^{26}\text{Mg}$ ,  $^{182}\text{Hf}$ - $^{182}\text{W}$ ,  $^{206}\text{Pb}$ - $^{207}\text{Pb}$ ) of CAIs, chondrules and achondrites suggest consistent relative ages that support the homogeneous distribution of  $^{26}\text{Al}$  throughout the disk and its chronolog-

ical significance. For example, the Allende CV3 chondrule ages determined from  $^{182}\text{Hf}$ - $^{182}\text{W}$  systematics ( $2.2 \pm 0.8$  Myr after CV CAIs [10]) and U-corrected Pb-Pb age ( $1.8 \pm 0.9$  Myr after CV CAIs [11]) are in good agreement with  $2.5$  ( $-0.4/+0.7$ ) Myr obtained from the  $(^{26}\text{Al}/^{27}\text{Al})_0$  of Kaba CV3 chondrules [7]. In contrast these are largely inconsistent with the old Pb-Pb ages of Allende chondrules [12], that might be due to the Pb-Pb ages compromised by the common Pb [13].

On the other hand, the Al-Mg and U-corrected Pb-Pb ages of volcanic angrites determined by Schiller et al. (2015) indicate  $^{26}\text{Al}$ - $^{26}\text{Mg}$  and  $^{207}\text{Pb}$ - $^{206}\text{Pb}$  ages compared to CV CAIs are inconsistent by  $\sim 1.5$  Myr. The U-corrected Pb-Pb ages of individual chondrules range from 0 to  $\sim 3\text{--}4$  Myr after CAIs, and their  $(^{26}\text{Al}/^{27}\text{Al})_0$  are much lower than those expected from their Pb-Pb ages, supporting the reduced abundance of  $^{26}\text{Al}$  in the disk regions where chondrules originated [12,14].

Because of this controversy, it is not clear what the differences in  $(^{26}\text{Al}/^{27}\text{Al})_0$  among and within chondrite groups represent: spatial heterogeneity of  $^{26}\text{Al}$  abundances in chondrule-forming region(s) and/or reflect multiple generations of chondrules formed at different times.

**Implication for thermal history of parent asteroids:** Despite the possibility of heterogeneous distribution of  $^{26}\text{Al}$  in the disk, the  $(^{26}\text{Al}/^{27}\text{Al})_0$  recorded in chondrules potentially provide an upper limit on  $^{26}\text{Al}$  abundances available as a heat source of their parent asteroids due to decay of  $^{26}\text{Al}$ . The inferred  $(^{26}\text{Al}/^{27}\text{Al})_0$  of chondrules are similar to/slightly higher than those indicated from thermal modeling of their parent asteroids [15], suggesting rapid accretion of chondrules into their parent bodies after their formation. The  $(^{26}\text{Al}/^{27}\text{Al})_0$  in the Kaba chondrules is too low to melt a CV parent asteroid [7] and contradicts the existence of a molten core (e.g., [16]).

We will present more details at the workshop and in a chapter of the forthcoming book.

**References:** [1] Hutchison and Hutchison (1989) *Nature* **337**, 238–241. [2] Kita and Ushikubo (2012) *MAPS*, **47**, 1108–1119. [3] Nagashima et al. (2014) *Geochem. J.* **48**, 561–570. [4] Schrader et al. (2016) GCA, in press. [5] Tenner et al. (2013) *LPSC*, **47**, #2010. [6] Krot et al. (2014) *LPSC*, **45**, #2142. [7] Nagashima et al. (2016) GCA, in press. [8] Bizzarro et al. (2004) *Nature*, **431**, 275–278. [9] Luu et al. (2015) *PNAS*, **112**, 1298–1303. [10] Budde G. (2016) *PNAS*, **113**, 2886–2891. [11] Huyskens et al. (2016) *LPSC*, **47**, #2727. [12] Connelly et al. (2012) *Science*, **338**, 651–655. [13] Kita (2015) *MAPS*, suppl. #5360. [14] Bollard (2015) *MAPS*, suppl. #5211. [15] Sugiura and Fujiya (2014) *MAPS*, **49**, 772–787. [16] Elkins-Tanton et al. (2011) *EPSL*, **305**, 1–10.